

Module 13

Earthquake

EARTHQUAKE

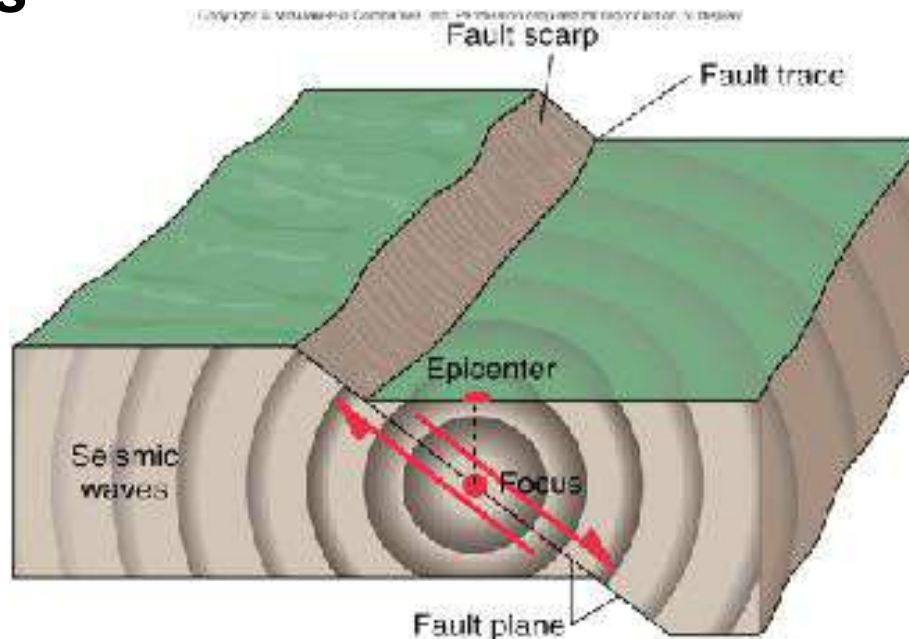


Photo credit: USGS

EARTHQUAKE

Earthquake

- ❑ A series of shock waves generated at a point (**the focus**) within the Earth's crust or mantle
- ❑ The point on the surface of the Earth above the focus is called the **epicenter**
- ❑ Three main types of wave motion are generated by an earthquake: P-Waves; S-Waves; dan L-Waves



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P-Waves

- High-frequency
- Short-wavelength
- Longitudinal waves
- Can be reflected and refracted
- Travel through the solid and the liquid part of the Earth

S-Waves

- High-frequency
- Short-wavelength
- Transverse waves
- Can be reflected and refracted
- Travel through the solid part of the Earth at varying velocities.
- Propagated in all directions from the focus

L-Waves

- Low-frequency
- Long-wavelength
- Transverse vibrations
- Confined to the outer skin of the crust
- Responsible for most of the destructive force of earthquake

EARTHQUAKE

Seismic Wave Types

Body Waves

- Primary or Compressional (**P-wave**)
- Secondary or Shear (**S-wave**)

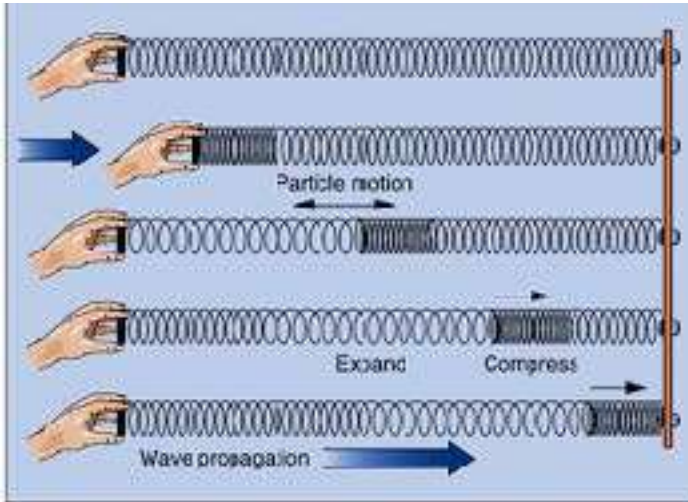
Surface Waves

- Rayleigh (large vertical displacements)
- Love (shear)

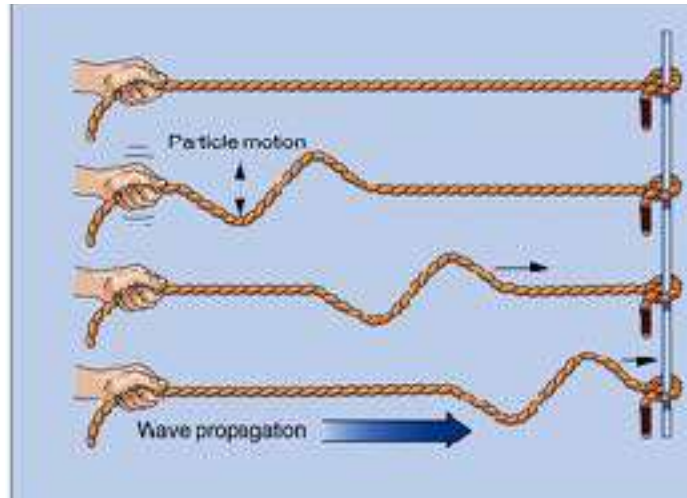
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Seismic Wave Types

Body Waves

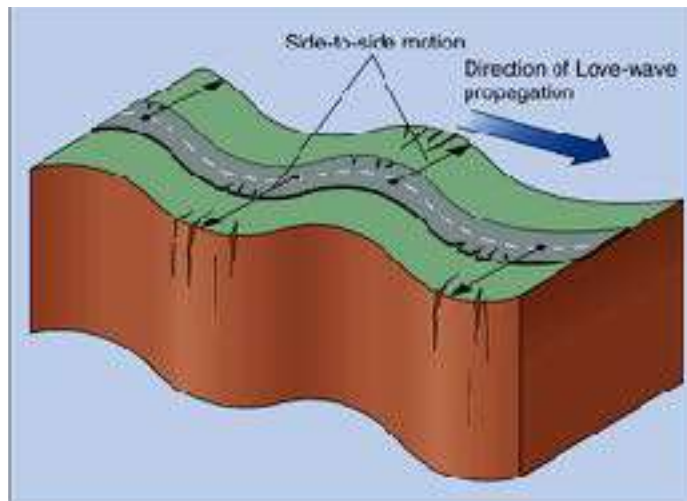


Primary (P) Wave

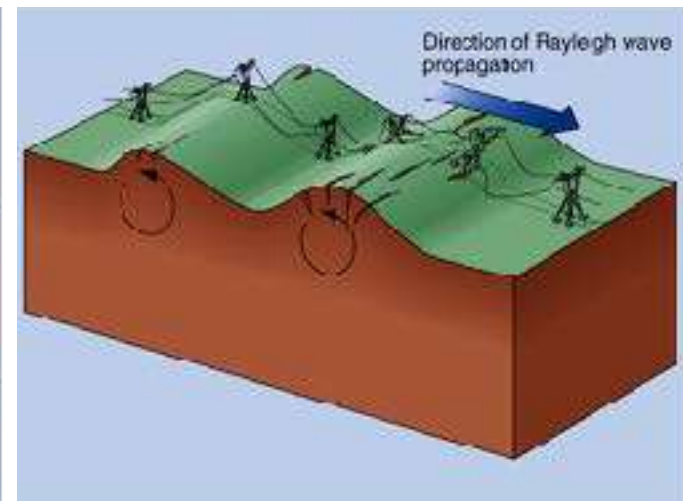


Secondary (S) Wave

Surface Waves



Love Wave

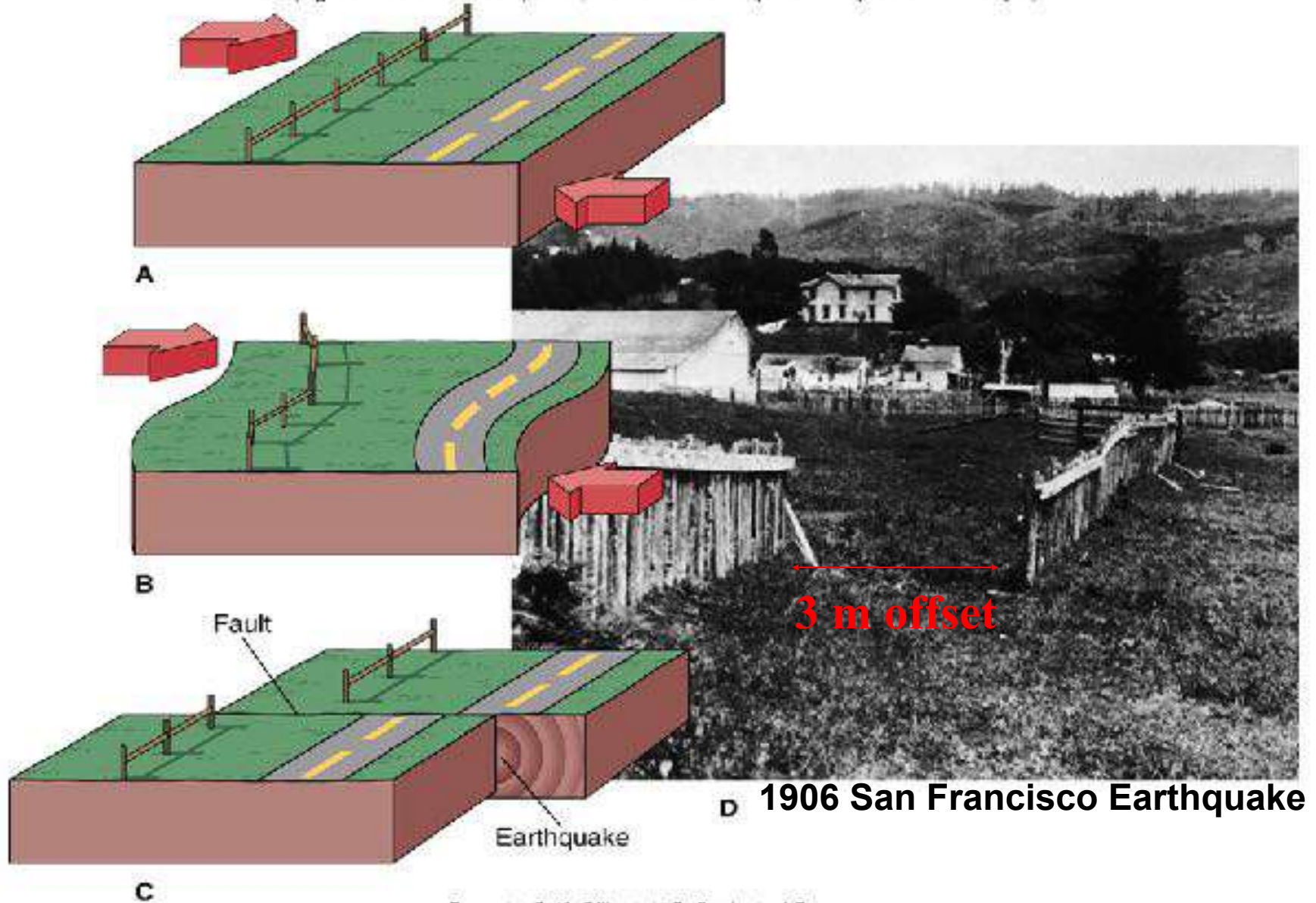


Rayleigh Wave

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Elastic Rebound Theory

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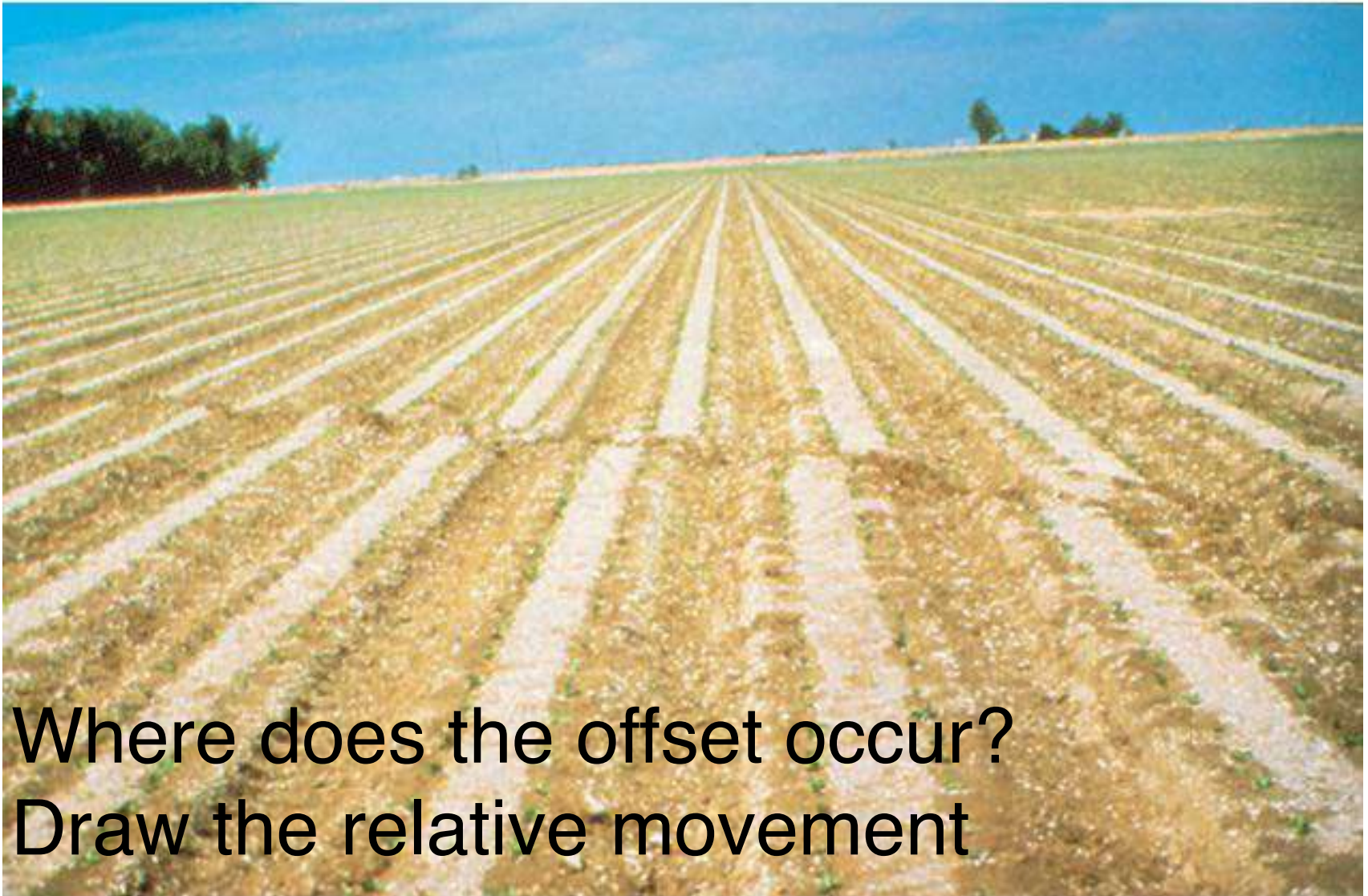


D 1906 San Francisco Earthquake

EARTHQUAKE

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Elastic Rebound Theory



Where does the offset occur?
Draw the relative movement

Offset Lettuce Rows - El Centro, CA

Photo by Univ. of Colorado; courtesy National Geographic Society, 1940

EARTHQUAKE

Equations for velocities

$$V_p = \left[\frac{k + 4/3\mu}{\rho} \right]^{1/2}$$

ρ density

μ shear modulus (rigidity)

k bulk modulus (rigidity)

$$V_s = \left[\frac{\mu}{\rho} \right]^{1/2}$$

because shear modulus (rigidity) for fluid is zero,
S waves cannot propagate through a fluid

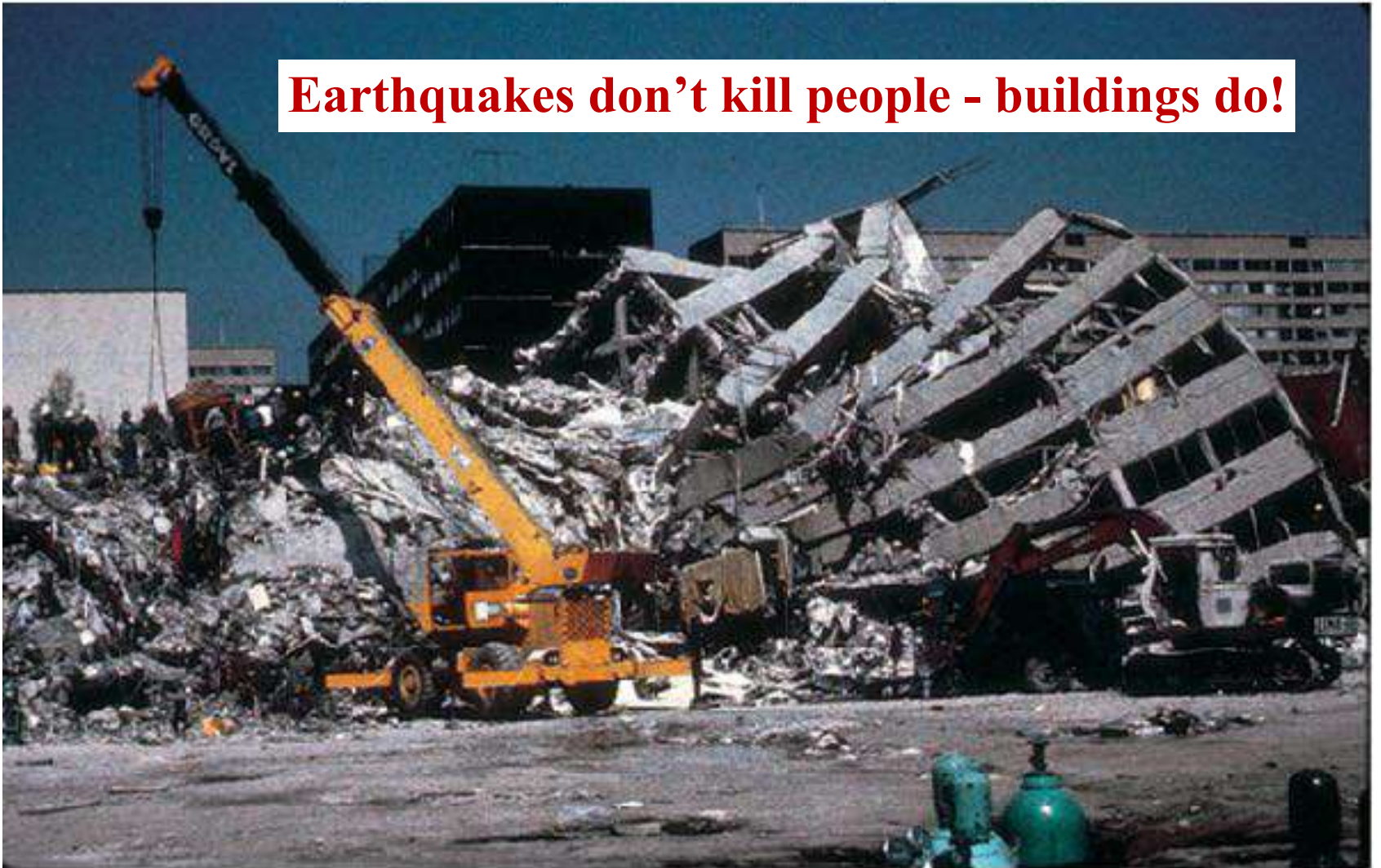
consequence of equations is that P-waves are 1.7x faster than S-waves

can infer physical properties from P and S waves

Effects of the Earthquake

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Earthquakes don't kill people - buildings do!



B

Pancaked Building - 1985 Mexico City

Photo by M. Celeli, U.S. Geological Survey

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B

Soil Liquefaction - 1964 Niigata, Japan

Photo by National Geophysical Data Center

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A

Surface Displacement - 1964 Alaska

Photo by U.S. Geological Survey

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B

Ground Rupture, 1906 Olema, CA

Photo by G. K. Gilbert, U.S. Geological Survey

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C

Fence Compression - Gallatin County, MT

Photo by I. J. Witkind, U.S. Geological Survey

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D

Buckled Concrete - 1971 San Fernando, CA

Photo by David McJeary

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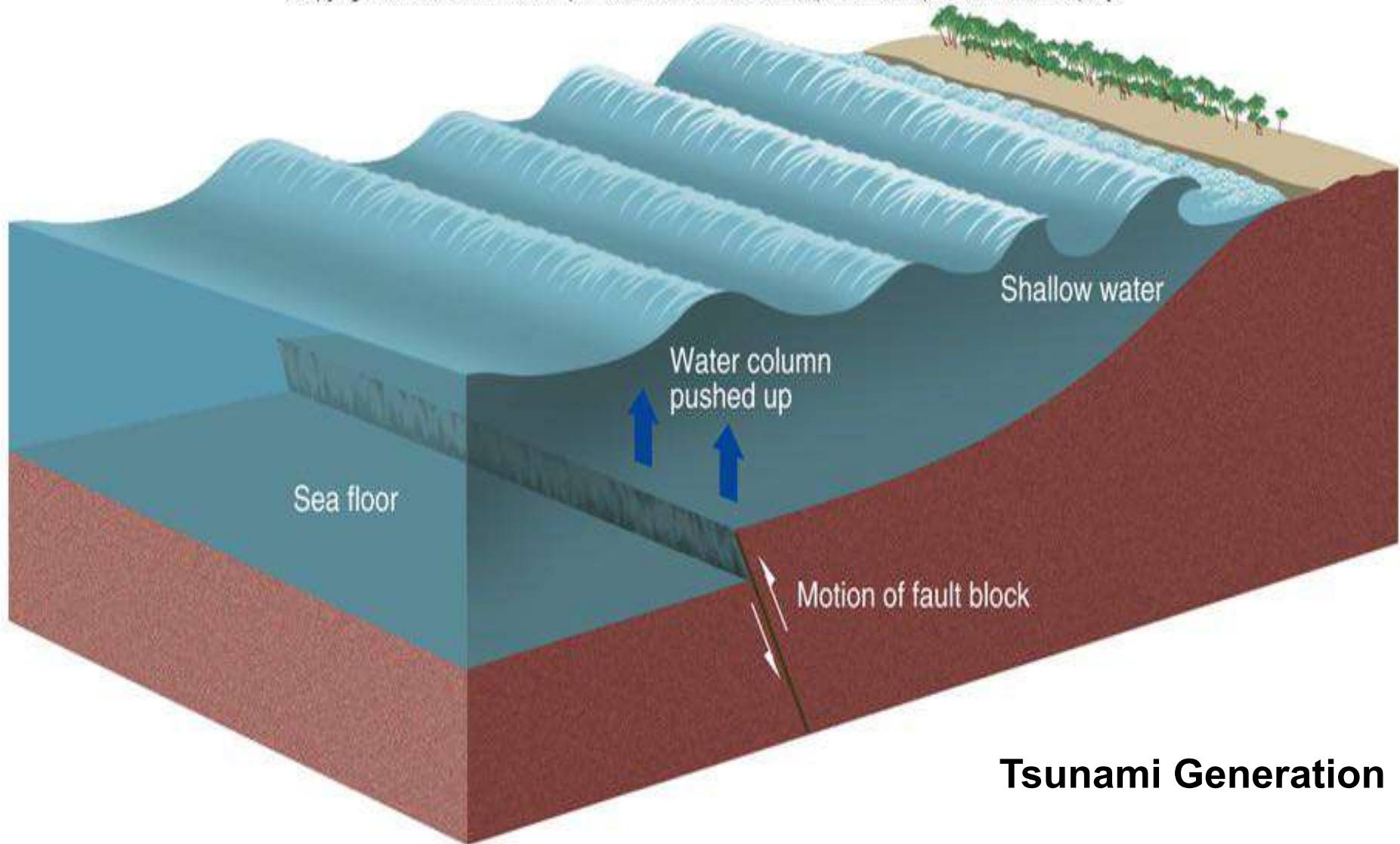
C

Tsunami Devastation - 1964 Alaska Earthquake

Photo by National Geophysical Data Center

Effects of the Earthquake

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Tsunami Generation

Effects of the Earthquake

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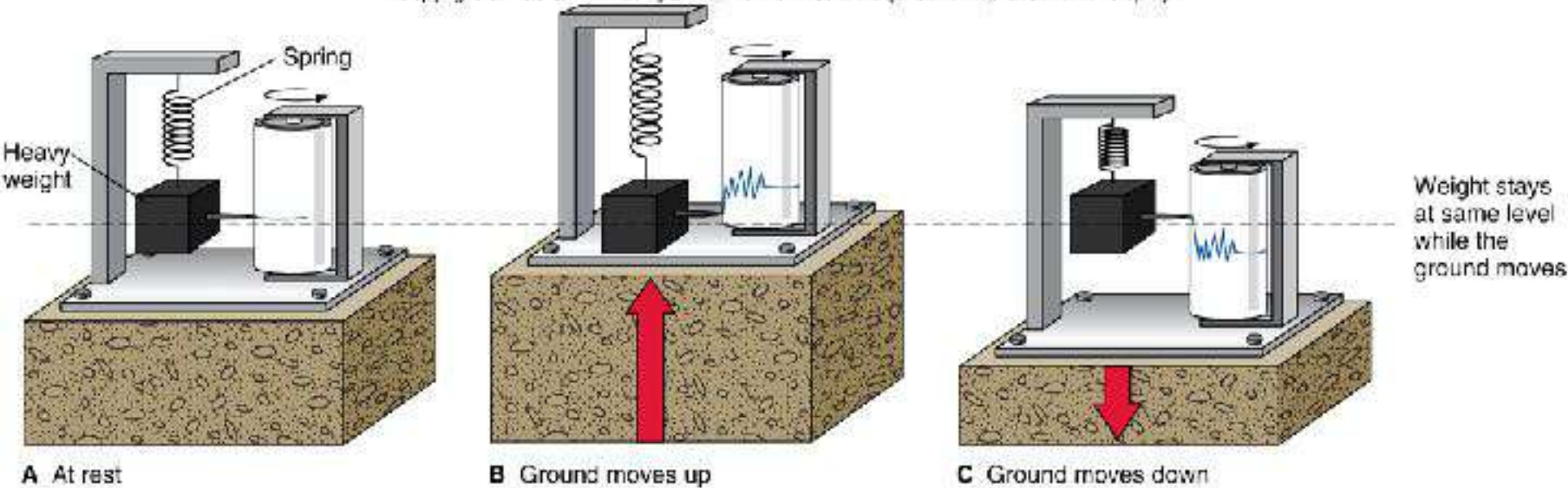


Tsunami Wave Propagation Times

Locating the Earthquake

Vertical Component Seismometer

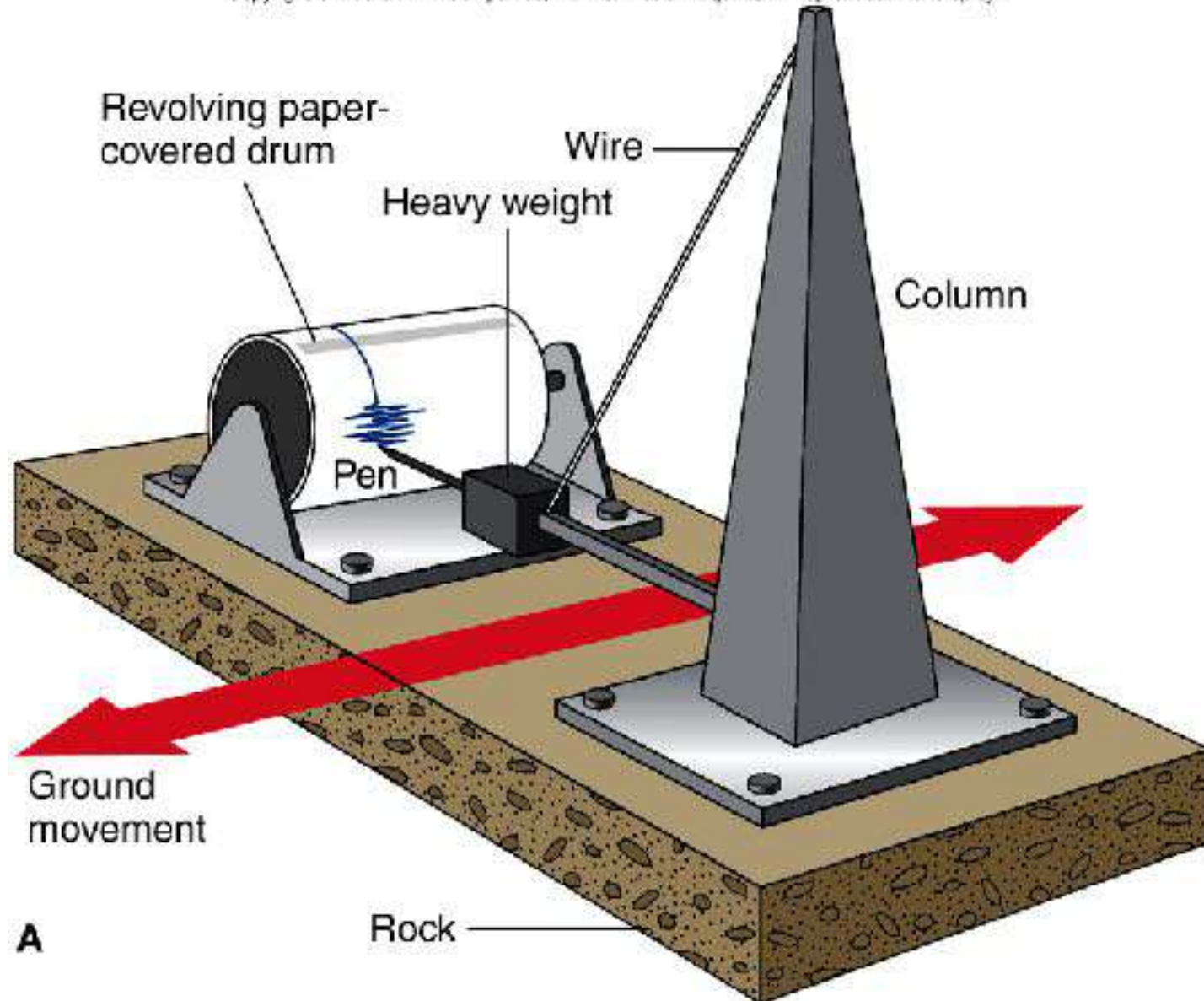
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Locating the Earthquake

Horizontal Component Seismometer

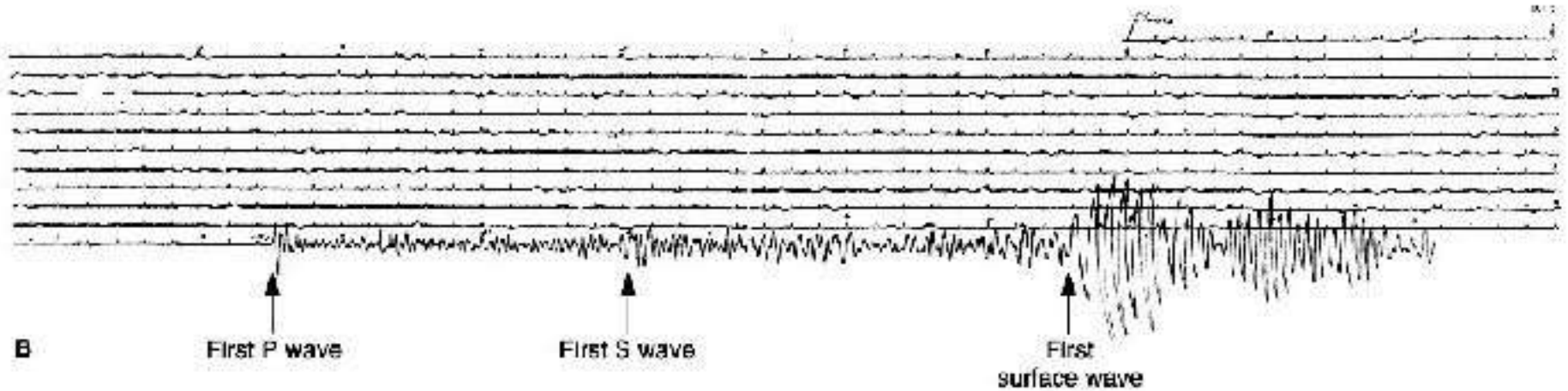
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Locating the Earthquake

First Arrivals – Seismographic Record

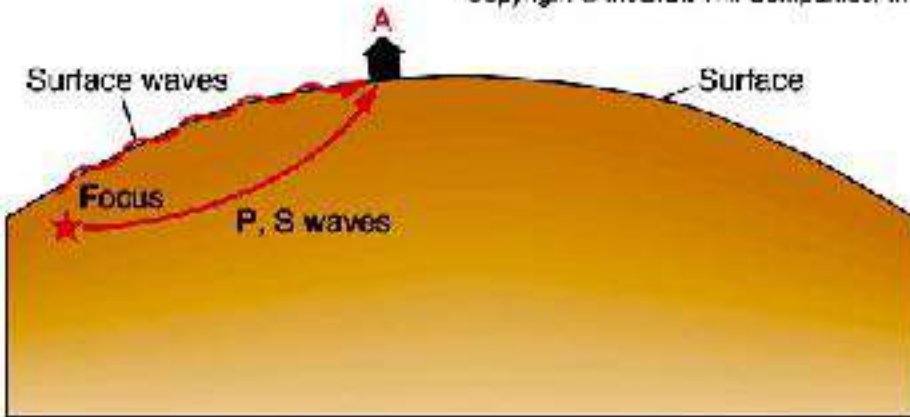
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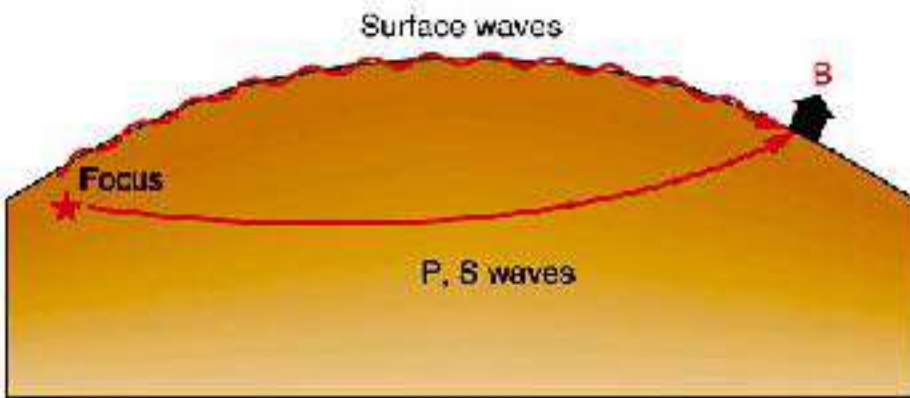
Locating the Earthquake

Distance – Time Relations

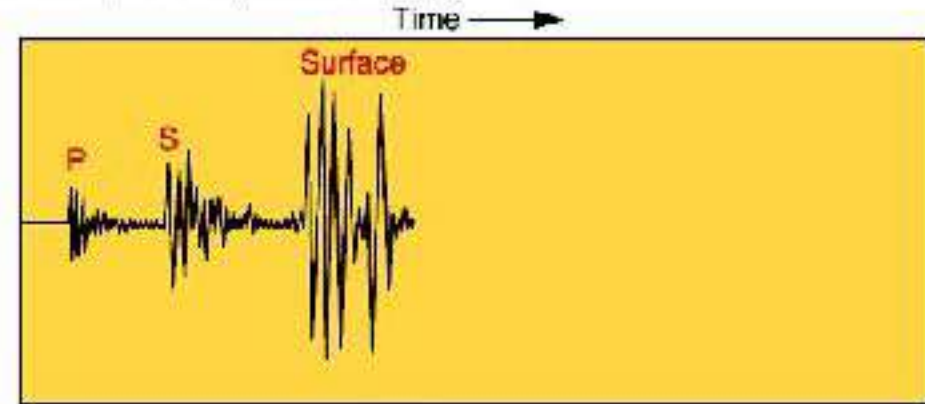
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A Station near focus

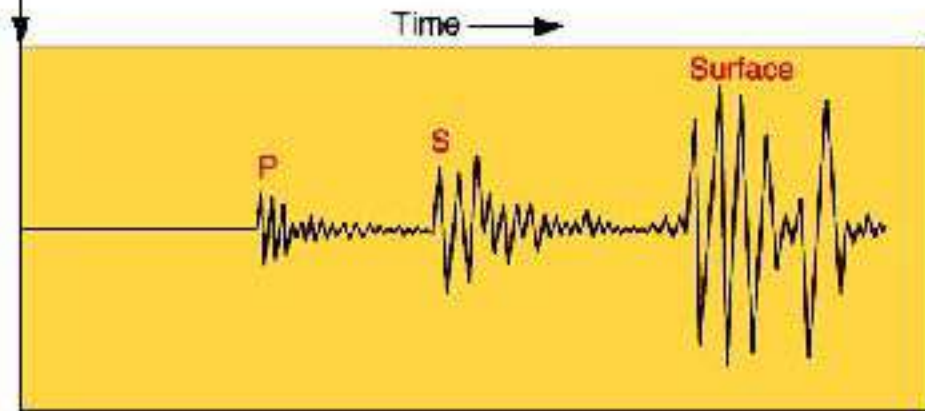


B Station far from focus



Seismogram from station A

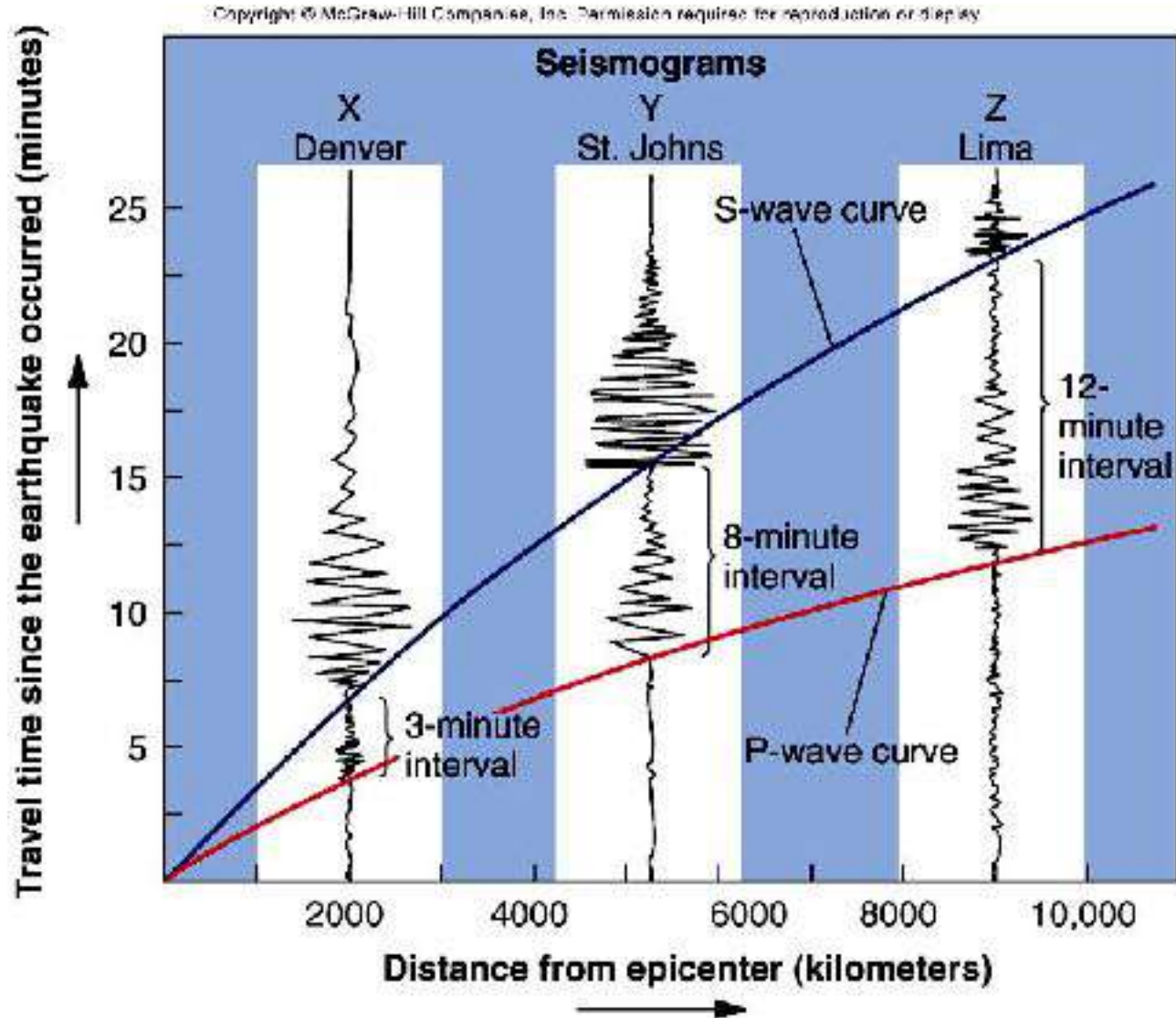
Time of earthquake



Seismogram from station B

Locating the Earthquake

P vs S Wave Travel Time Curves



Locating the Earthquake

Earthquake Location by Range



Measuring the Earthquake

- 3 distinct methods to measure earthquakes. Two based on energy and one based on intensity.
 - **Richter Magnitude Scale:** originally developed for southern California. Log scale, which has no upper bound. Small earthquakes may yield negative values. Tends to be inaccurate at >7 magnitudes.
 - **Moment Magnitude Scale:** measurement of the amount of work done during the earthquake. Based on rock strength, area of rupture, and displacement during event.
 - **Modified Mercalli Intensity Scale:** based on the damage associated with a particular event at a particular location. Ranges from I (less damage) to XII (most damage).

Measuring the Earthquake

Magnitude-Description-Intensity-Frequency Relations

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Table 16.2

Comparison of Earthquake Magnitude, Description, Intensity, and Expected Annual World Occurrence

Richter Magnitude	Description	Maximum Expected Mercalli Intensity at Epicenter	Annual Expected Number
2.0	Very Minor	I Usually detected only by instruments	600,000
2.0–2.9	Very Minor	I–II Felt by some indoors; especially on upper floors	300,000
3.0–3.9	Minor	III Felt indoors	49,000
4.0–4.9	Light	IV–V Felt by most; slight damage	6,200
5.0–5.9	Moderate	VI–VII Felt by all; damage minor to moderate	800
6.0–6.9	Strong	VII–VIII Everyone runs outdoors; moderate to major damage	266
7.0–7.9	Major	IX–X Major damage	18
8.0 or higher	Great	X–XII Major and total damage	1 or 2

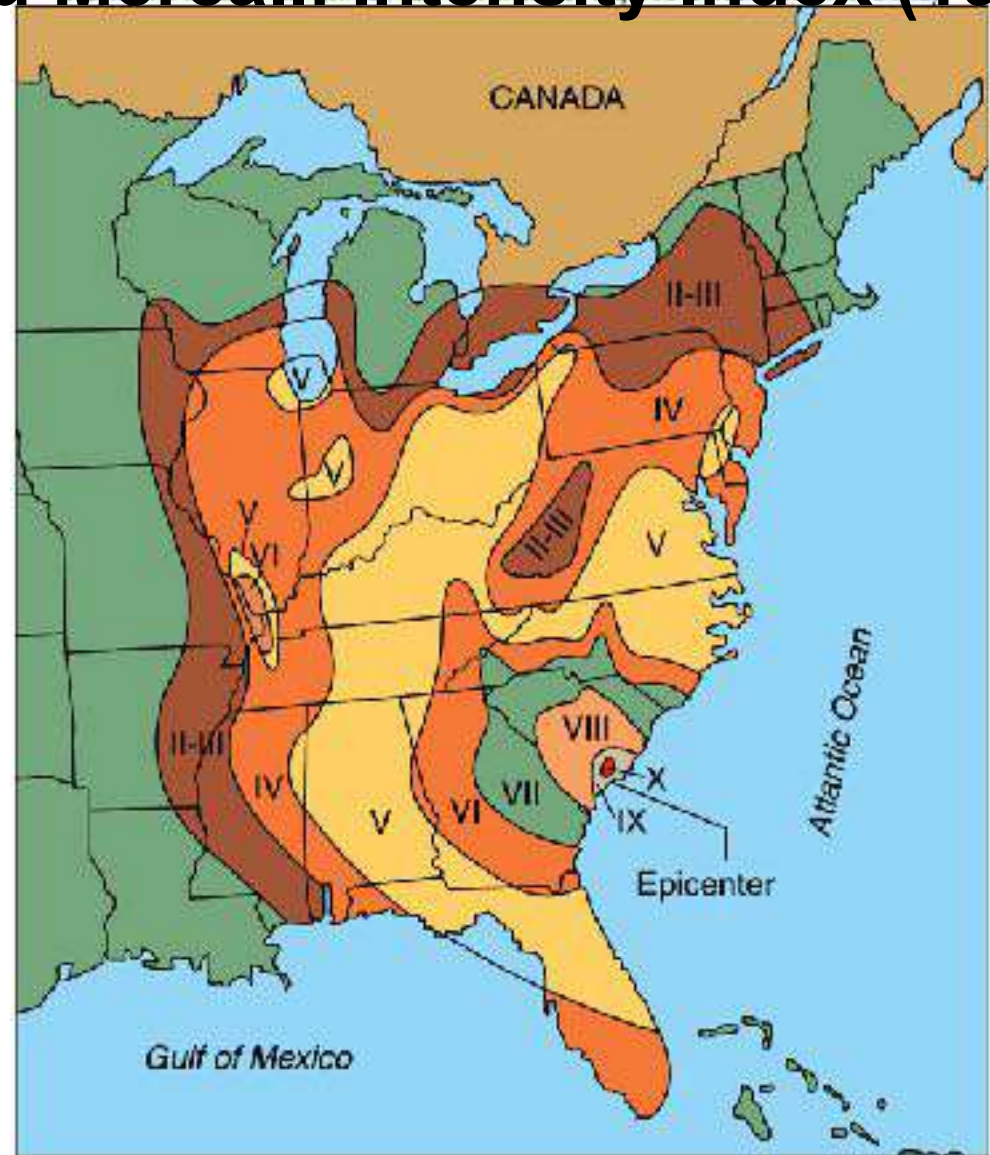
Source: U.S. Geological Survey

Measuring the Earthquake

Modified Mercalli Intensity Index (1931)

- I. Not felt except by a very few under especially favorable circumstances.
- II. Felt only by a few persons at rest, especially on upper floors of buildings. Delicately suspended objects may swing.
- III. Felt quite noticeably indoors, especially on upper floors of buildings but many people do not recognize it as an earthquake. Standing motor cars may rock slightly. Vibration like passing of truck. Duration estimated.
- IV. During the day felt indoors by many, outdoors by few. At night some awakened. Dishes, windows, doors disturbed; walls made cracking sound. Sensation like heavy truck striking building. Standing motor cars rocked noticeably.
- V. Felt by nearly everyone; many awakened. Some dishes, windows, etc., broken; a few instances of cracked plaster; unstable objects overturned. Disturbance of trees, poles, and other tall objects sometimes noticed. Pendulum clocks may stop.
- VI. Felt by all; many frightened and run outdoors. Some heavy furniture moved; a few instances of fallen plaster or damaged chimneys. Damage slight.
- VII. Everybody runs outdoors. Damage negligible in buildings of good design and construction; slight to moderate in well-built ordinary structures; considerable in poorly built or badly designed structures; some chimneys broken. Noticed by persons driving motor cars.
- VIII. Damage slight in specially designed structures; considerable in ordinary substantial buildings with partial collapse; great in poorly built structures. Panel walls thrown out of frame structures. Fall of chimneys, factory stacks, columns, monuments, walls. Heavy furniture overturned. Glass and mud ejected in small amounts. Changes in well water. Persons driving motor cars disturbed.
- IX. Damage considerable in specially designed structures; well-designed frame structures thrown out of plumb; great in substantial buildings, with partial collapse. Buildings shifted off foundations. Ground cracked conspicuously. Underground pipes broken.
- X. Some well-built wooden structures destroyed; most masonry and frame structures destroyed with foundations; ground badly cracked. Rails bent. Considerable landslides from river banks and steep slopes. Shifted sand and mud. Water splashed (slopped) over banks.
- XI. Few if any, (masonry) structures remain standing. Bridges destroyed. Broad fissures in ground. Underground pipelines completely out of service. Earth slumps and land slips in soft ground. Rails bent greatly.
- XII. Damage total. Waves seen on ground surface. Lines of sight and level distorted. Objects thrown upward into the air.

From Wood and Neumann, 1931, Bulletin Seismological Society of America



1886 Charleston, SC earthquake

Measuring the Earthquake

Historical Earthquake Magnitudes

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Table 16.3 Earthquake Magnitudes

		Richter Magnitude	Moment Magnitude
1811-12	New Madrid, Missouri area	7.5, 7.3, 7.6	7.7, 7.6, 7.9
1867	Fert Tojco, S. Calif.	7.6	7.6
1872	Lane Pine, Calif.	7.3	7.6
1886	Charleston, South Carolina	6.7	7.0
1906	San Francisco, N. Calif.	8.25	7.7
1915	Heaven Valley, Nevada	7.7	7.1
1933	Long Beach, S. Calif.	6.3	6.2
1962	San Geronimo, S. Calif.	7.2	7.5
1964	Dixie Valley, Nevada	7.2, 7.1	7.5, 6.9
1957	Aloufan Islands, Alaska	8.1	8.8
1969	Southeastern Alaska	7.9	8.5
1959	Hebgen Lake, Montana	7.7	7.5
1960	Chile	8.5	9.0
1964	Great Anchorage, Alaska	6.0	6.2
1965	Aloufan Islands, Alaska	8.2	8.7
1970	Honi	7.75	7.9
1971	San Fernando Valley, S. Calif.	6.4	6.6
1975	Hawaii	7.2	7.5
1976	China	7.6	7.6
1980	Lambert County, N. Calif.	6.9	7.2
1983	Coalinga, Calif.	6.7	6.2
1983	Challis, Idaho	7.2	7.0
1989	Adirondack Mountains, New York	6.1	4.6
1989	Hawaii	6.5	
1985	Atlix, Mexico		8.1, 7.5
1987	Lynnwoodville, Illinois	5.0	5.0
1987	Whetser, S. Calif.	5.9	5.0
1988	Quebec	5.0	
1989	Loma Prieta, N. Calif.	7.0	7.2
1989	Hawaii	6.1	6.4
1982	Lambert County, N. Calif.	7.1, 6.6, 6.7	
1992	Landers, S. Calif.	7.8, 6.7	7.56, 6.4
1994	Northridge, N. Calif.	6.4	6.7
1995	Kobe, Japan (7.2 on Japanese scale)	6.5	6.8
1998	Papua New Guinea		7.1
1999	Amrit, Turkey		7.4
1999	Taipei, Taiwan		7.6
2001	El Salvador		7.7
2001	Bujarat, India		7.7
2001	Puget Sound (Nisqually), Washington		6.8

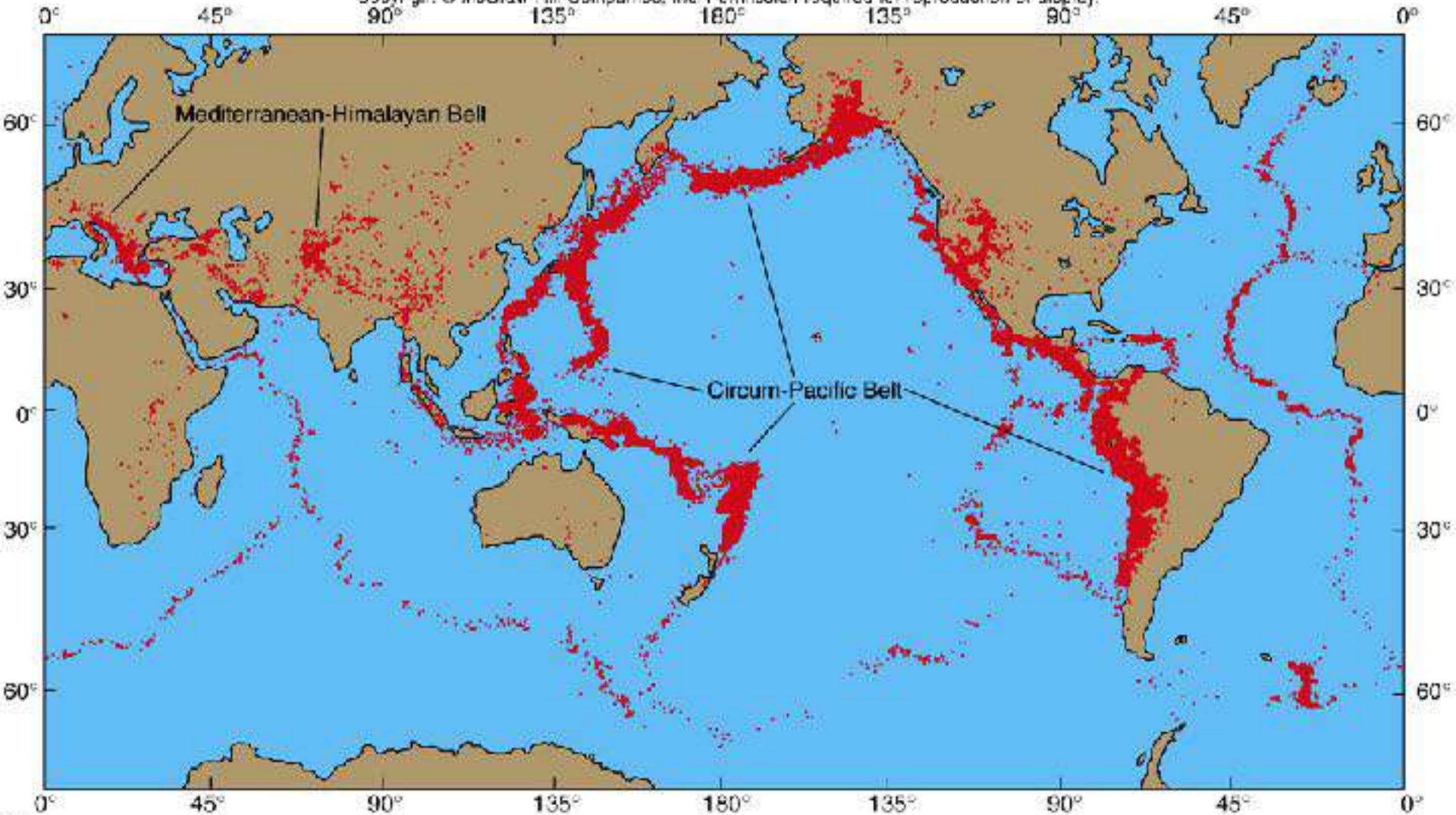
Earthquake and Plate Tectonics

- Earthquakes at Plate Boundaries
- Subduction Angle

Earthquake and Plate Tectonics

Earthquake Distribution

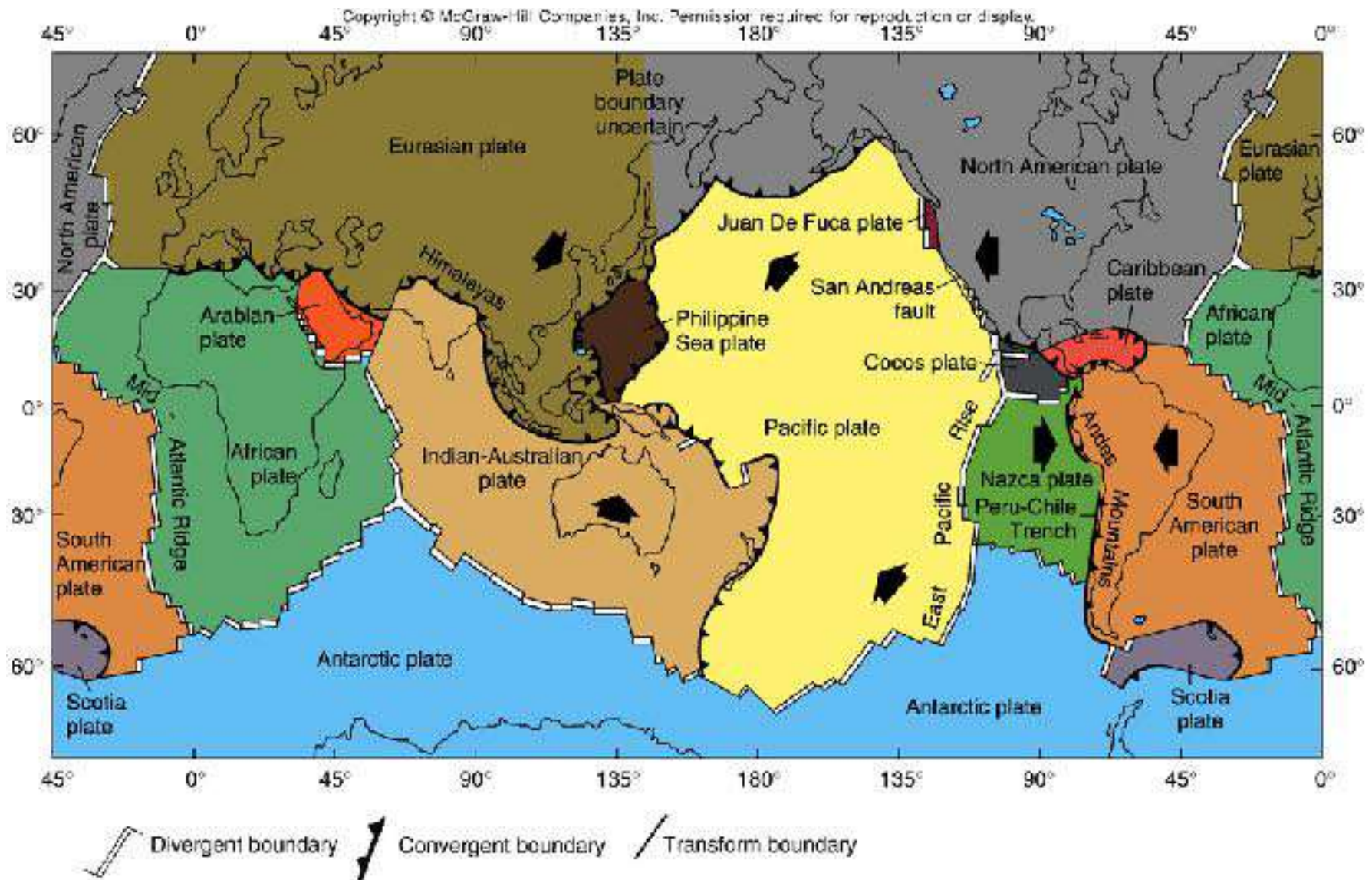
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A

Earthquake and Plate Tectonics

Relative plate motion and boundaries

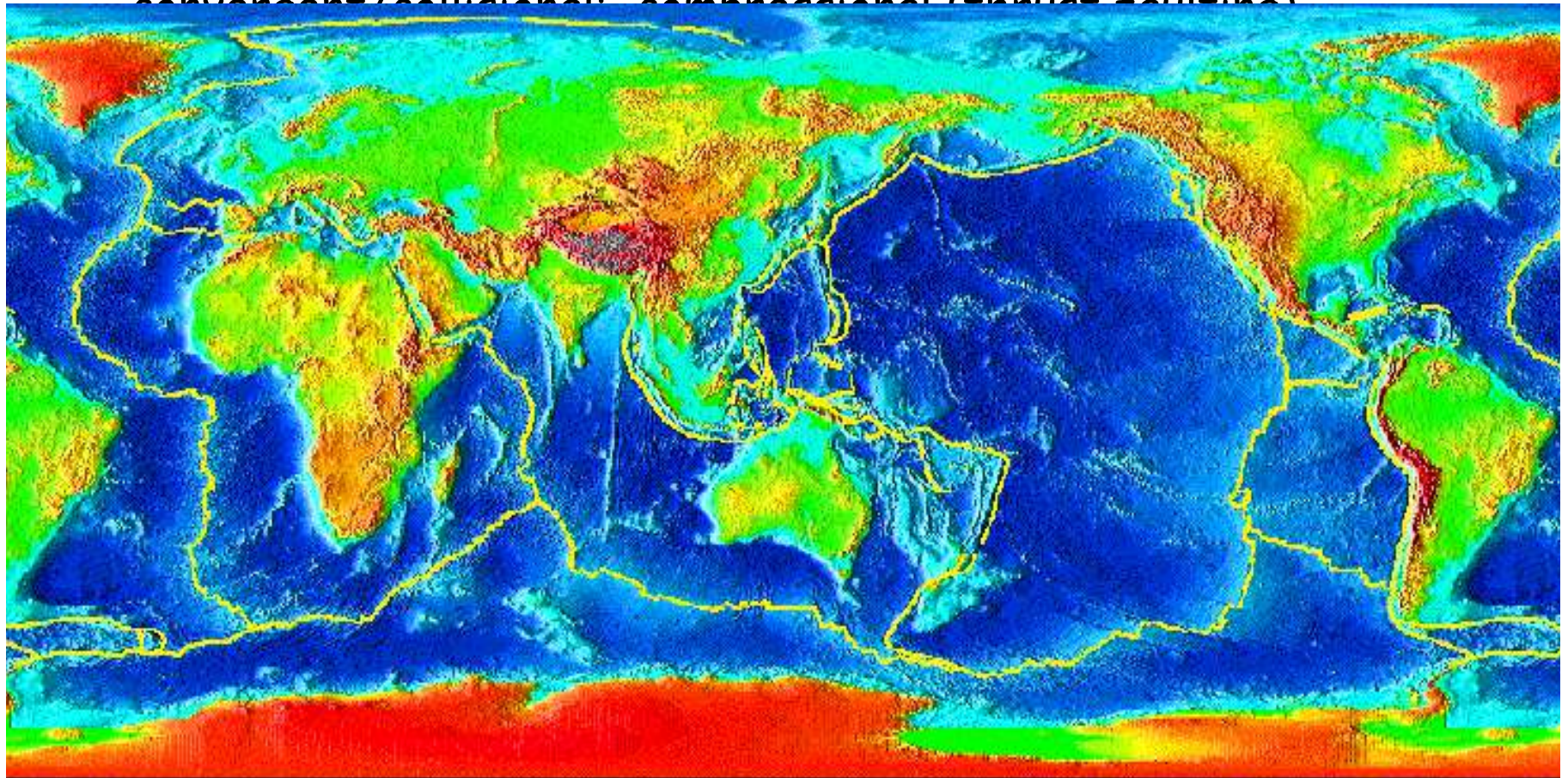


Earthquake and Plate Tectonics

different types of structures are associated with each boundary type:

divergent/rifting: extensional (normal faulting)

convergent/collisional: compressional (thrust faulting)

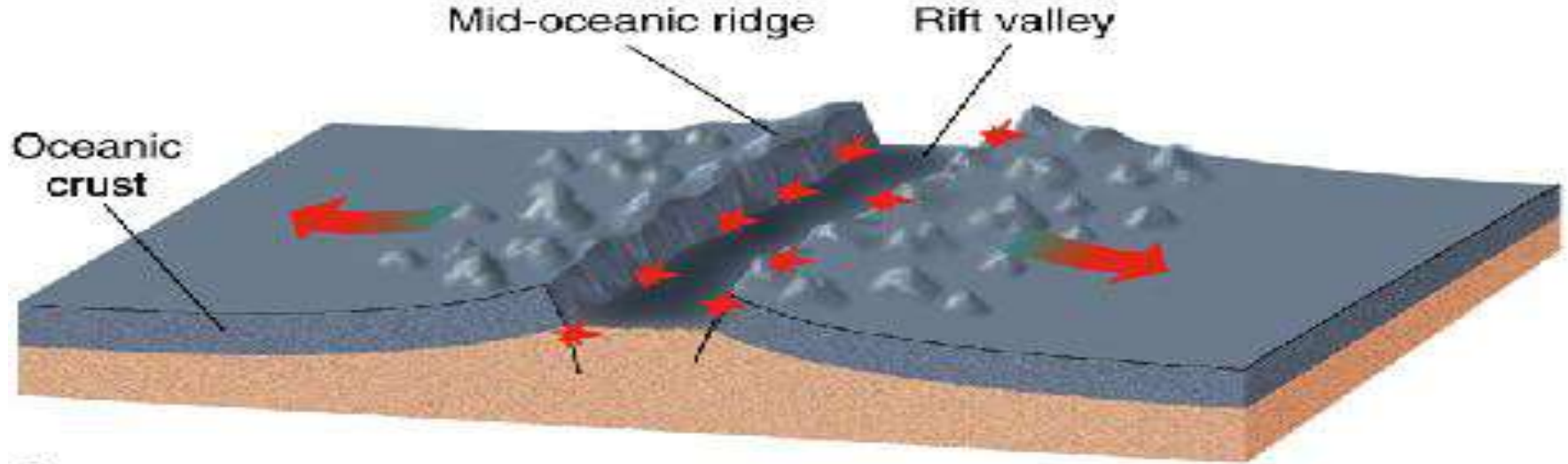


Crustal Plate Boundaries

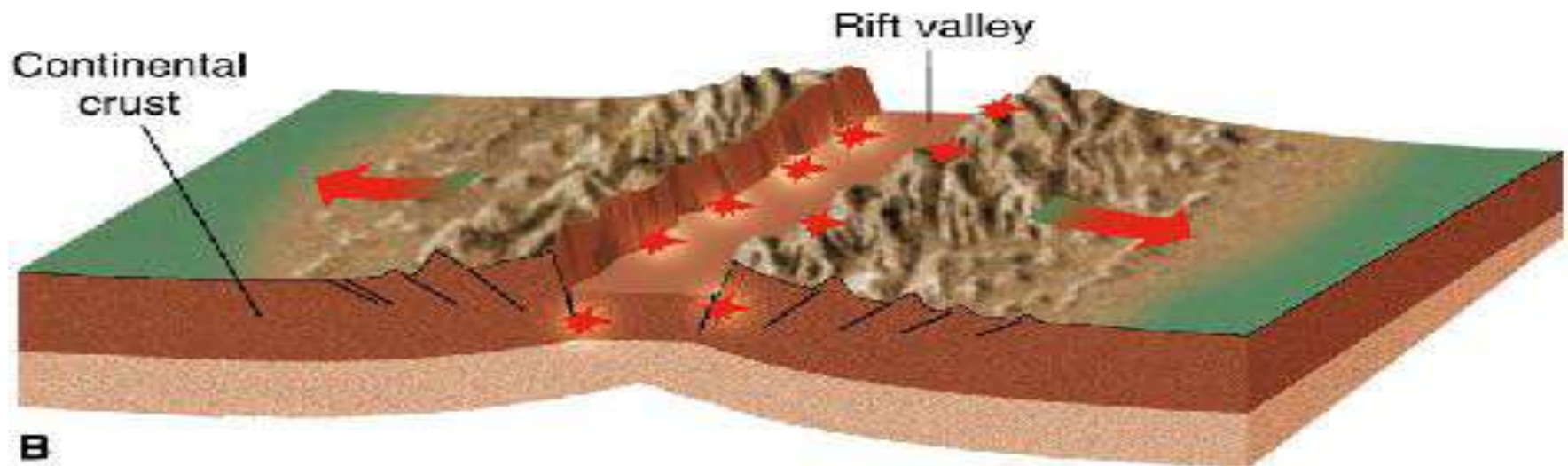
Earthquake and Plate Tectonics

Divergent Boundary and Earthquakes

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A

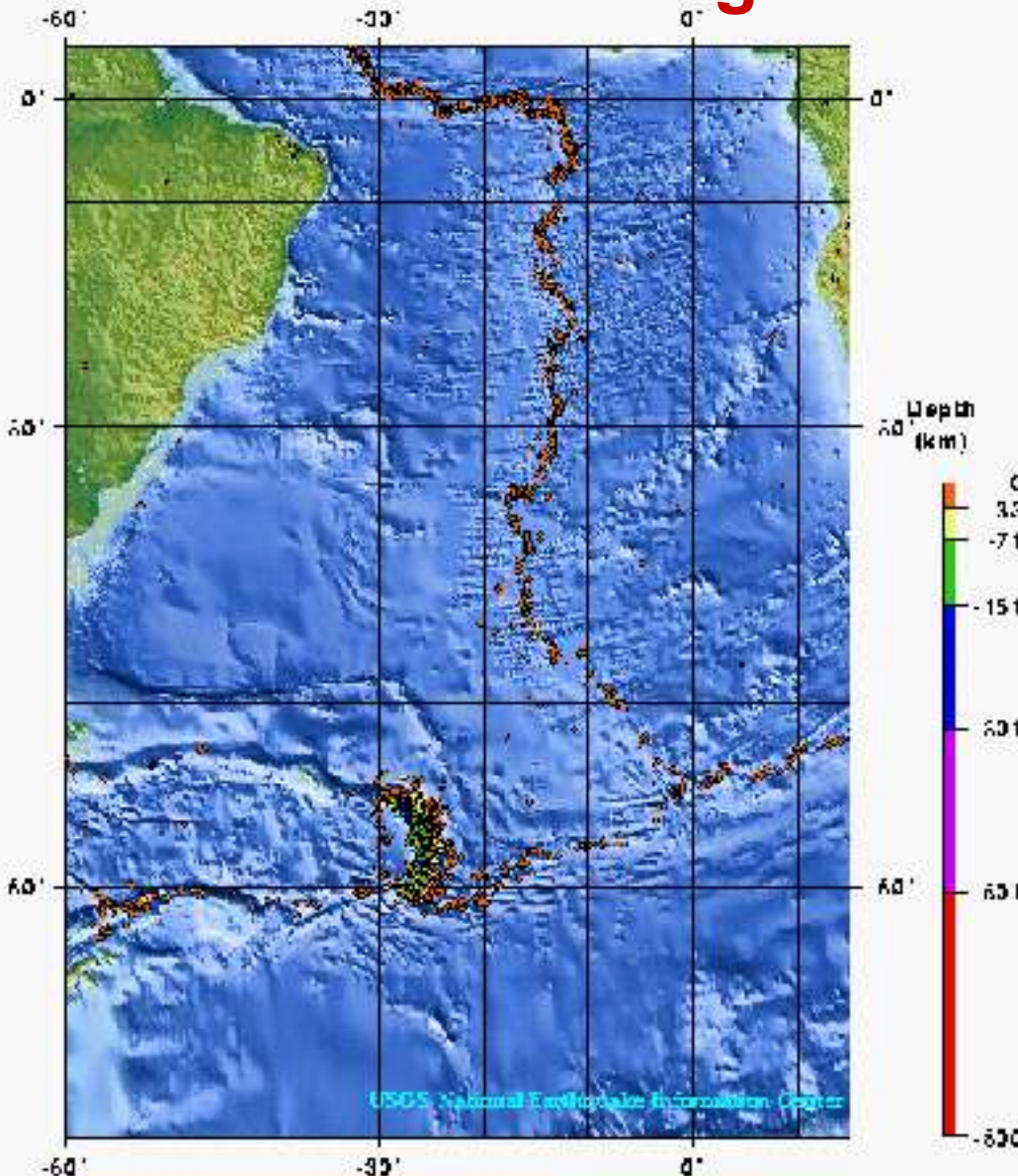


B

Earthquake and Plate Tectonics

Seismicity of the South Atlantic Ocean: 1975-1995

Divergent Boundary and Earthquakes

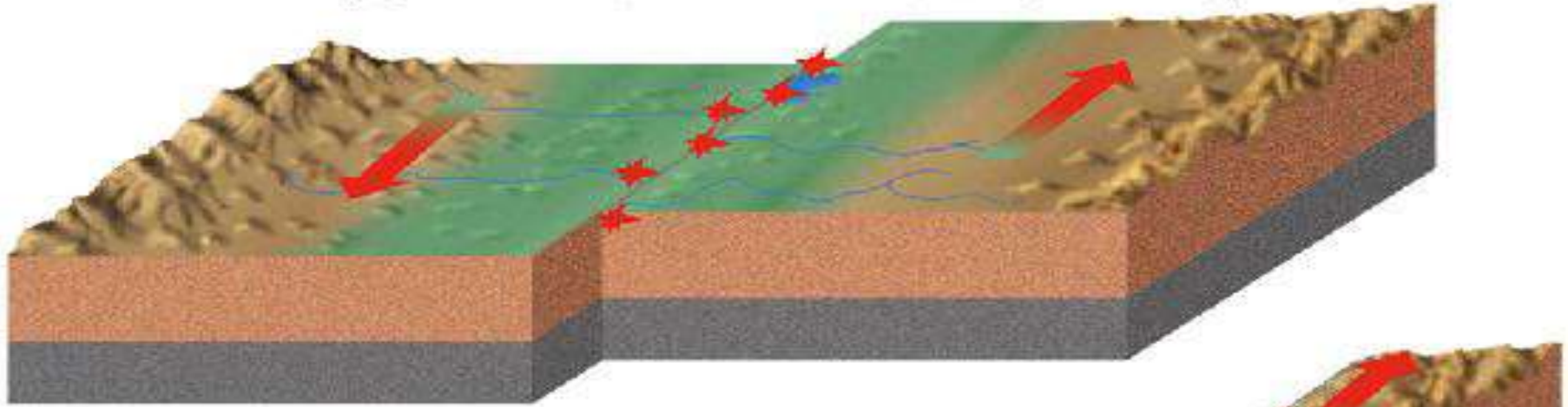


Earthquakes along mid-ocean ridge are shallower than those along subduction zone

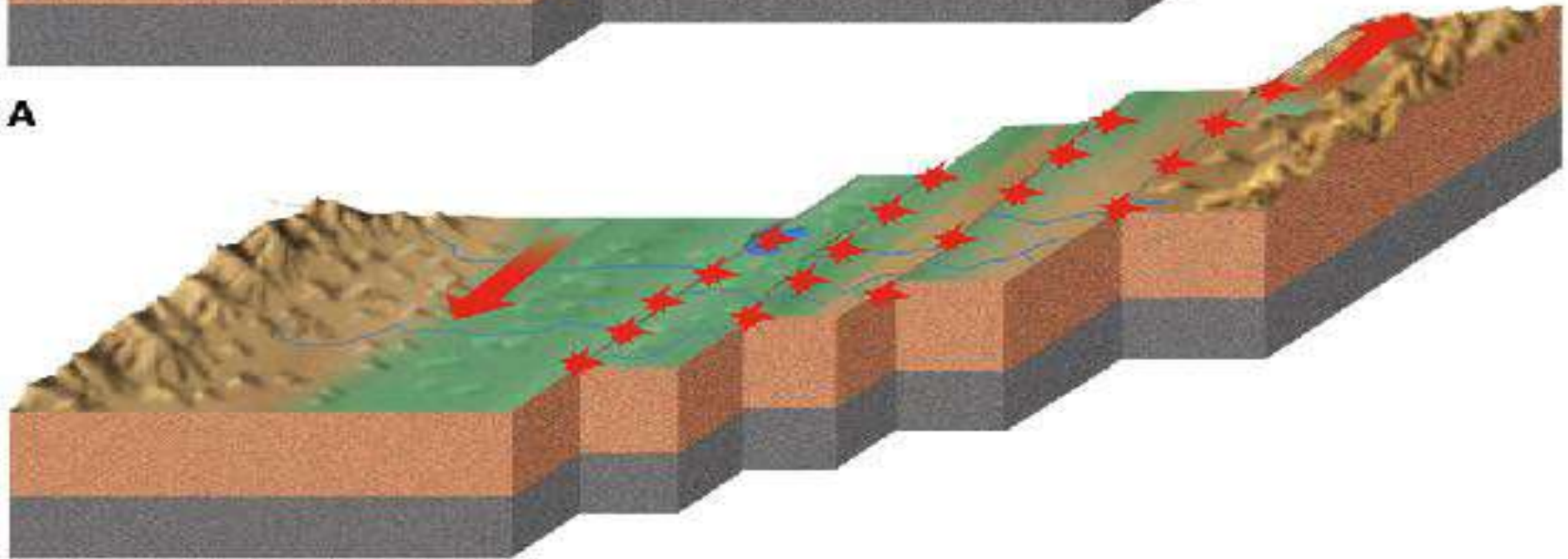
Earthquake and Plate Tectonics

Transform Boundary and Earthquakes

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A

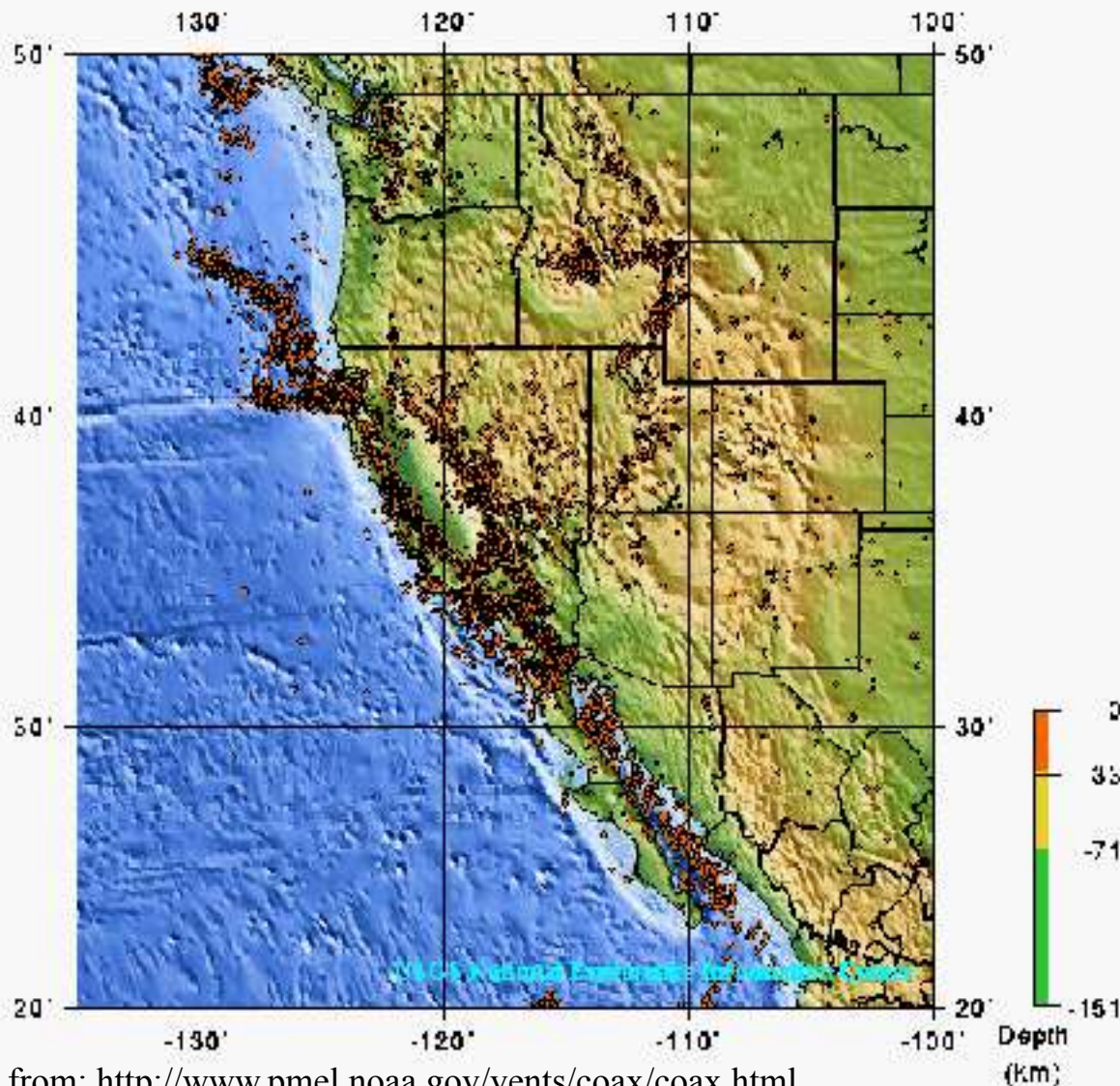


B

Earthquake and Plate Tectonics

Sismicity of the Western United States, 1973-1993

Transform Boundary and Earthquakes



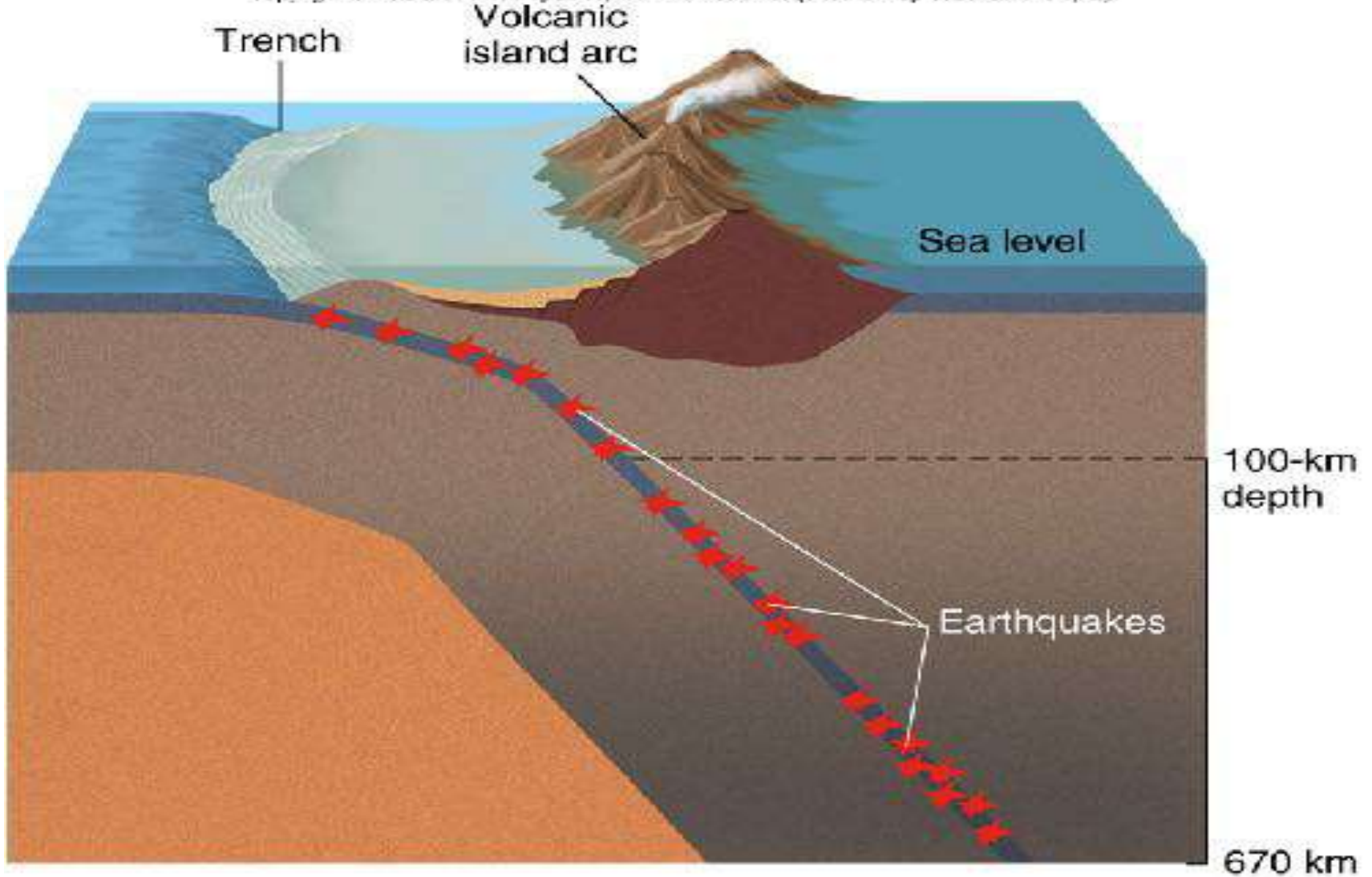
**The western US
is somewhat
anomalous
Note: absence of
deep earthquakes**

from: <http://www.pmel.noaa.gov/vents/coax/coax.html>

Earthquake and Plate Tectonics

Convergent Boundary and Earthquakes

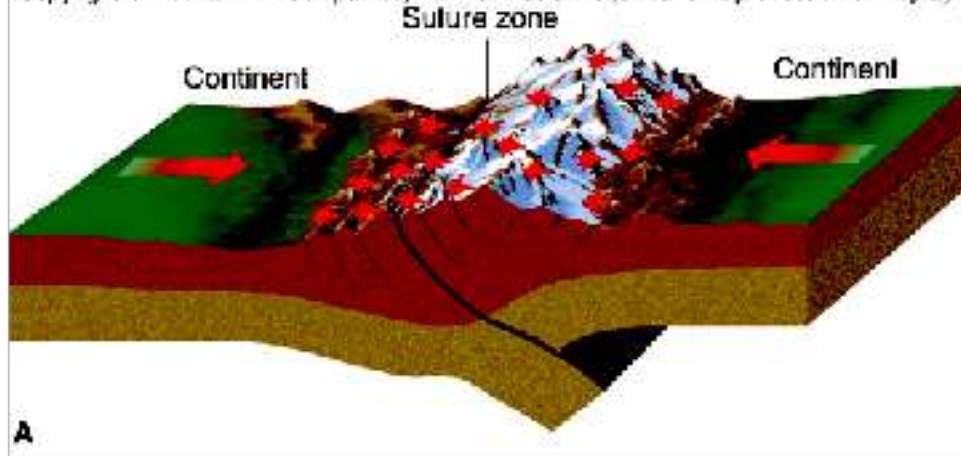
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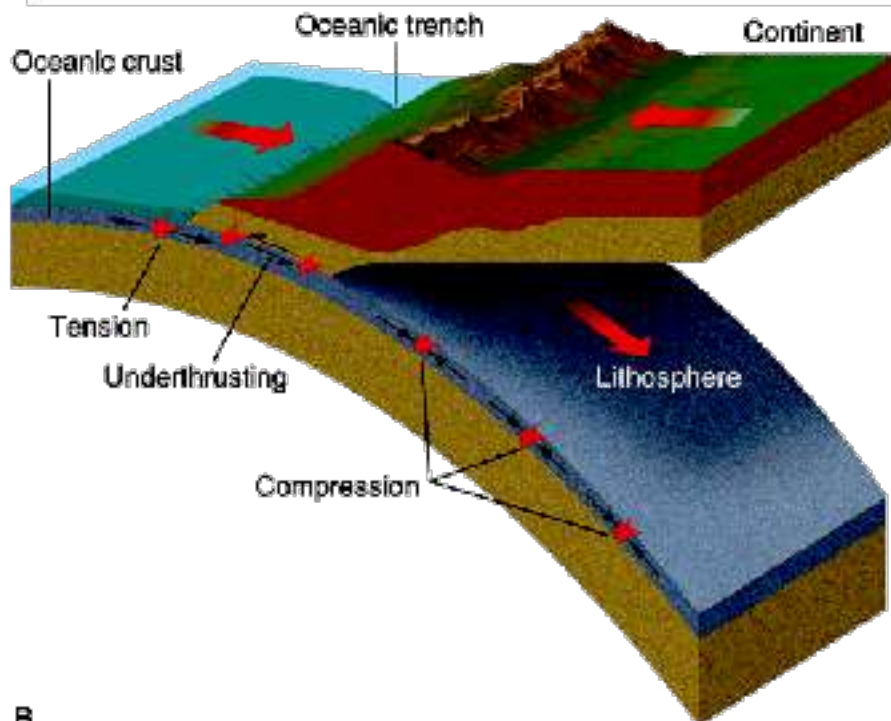
Earthquake and Plate Tectonics

Shallow vs. Deep Subduction Earthquakes

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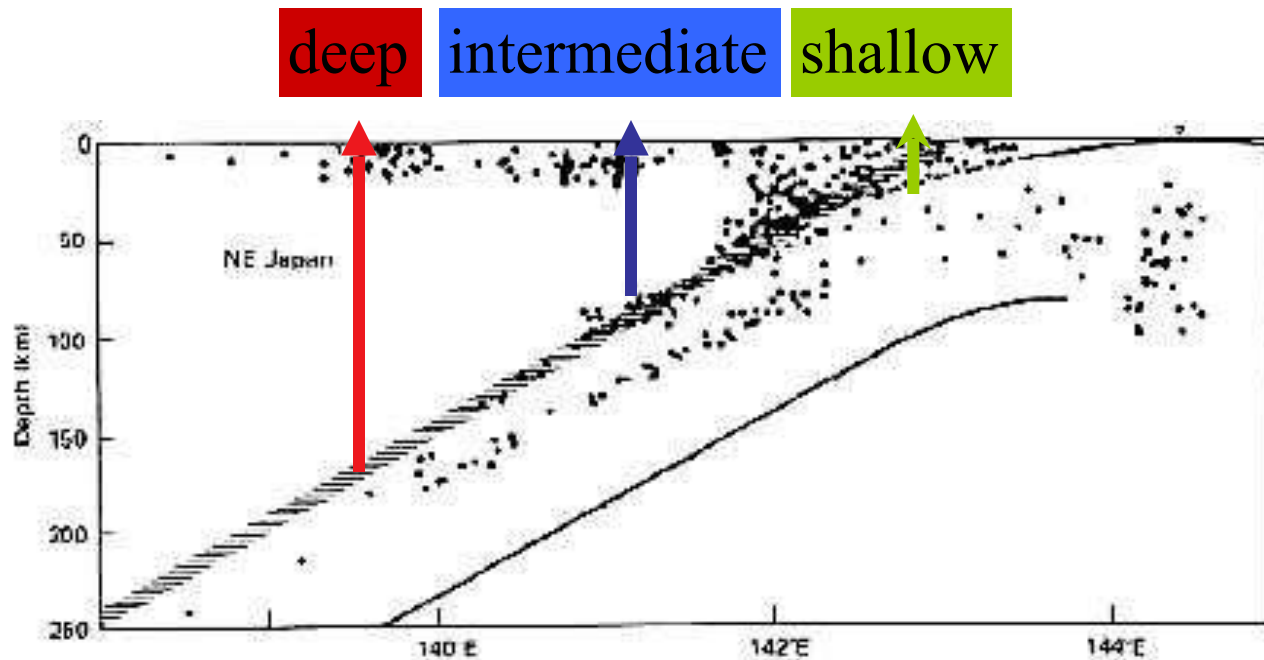


Continent-continent collision zones have broad areas of relatively shallow seismicity



Ocean-continent convergent margins have earthquake foci that extend to great depths. Mechanism tend to change from extension to compression downdip.

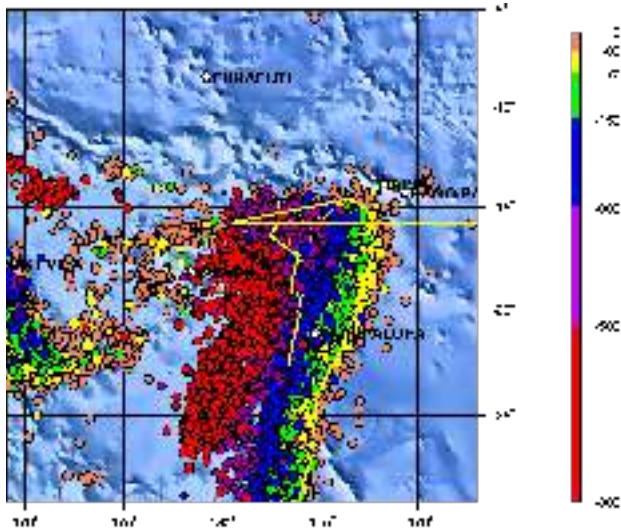
Earthquake and Plate Tectonics



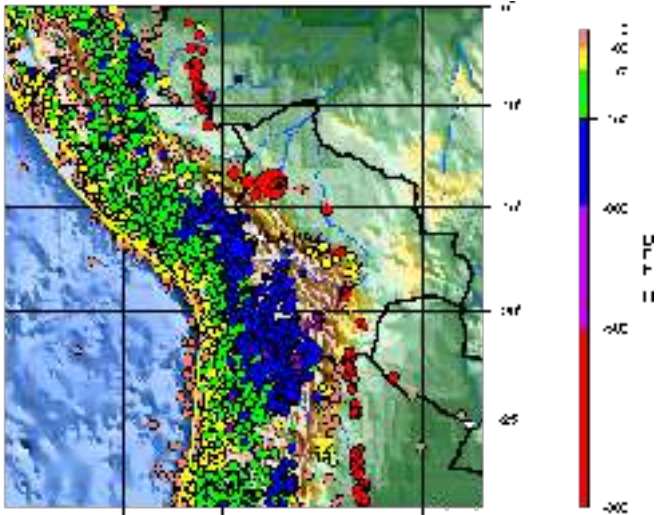
- **Epicenters: location of earthquake rupture projected to surface;**
- **Dip of slab leads to observed seismicity patterns: deeper farther from trench**
- **Location of downgoing slab as it dives into mantle is defined by seismicity.**
- **Earthquakes occur along an inclined belt: the Wadati-Benioff zone reaches maximum depth of ~670 km**

Earthquake and Plate Tectonics

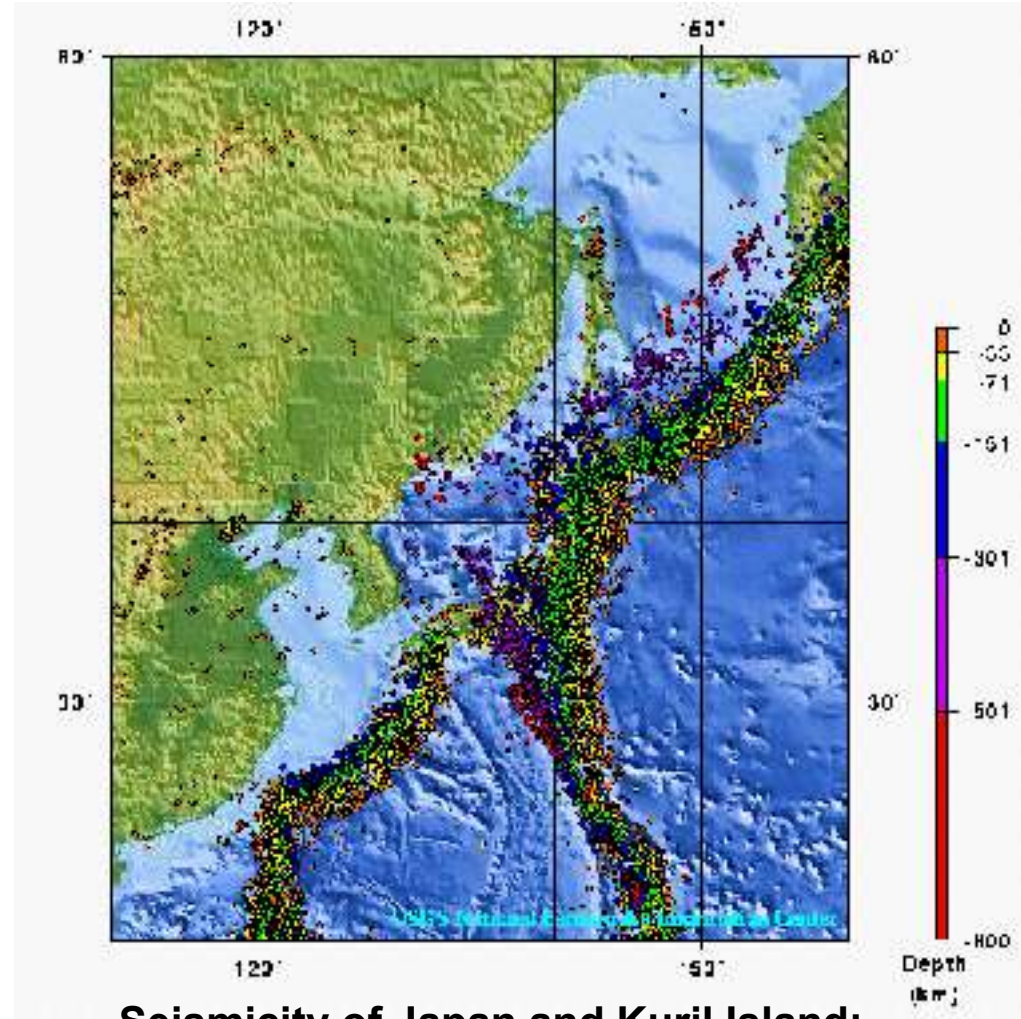
Seismicity of subduction zones



**Seismicity Fiji Islands
Region: 1977 - 1997**



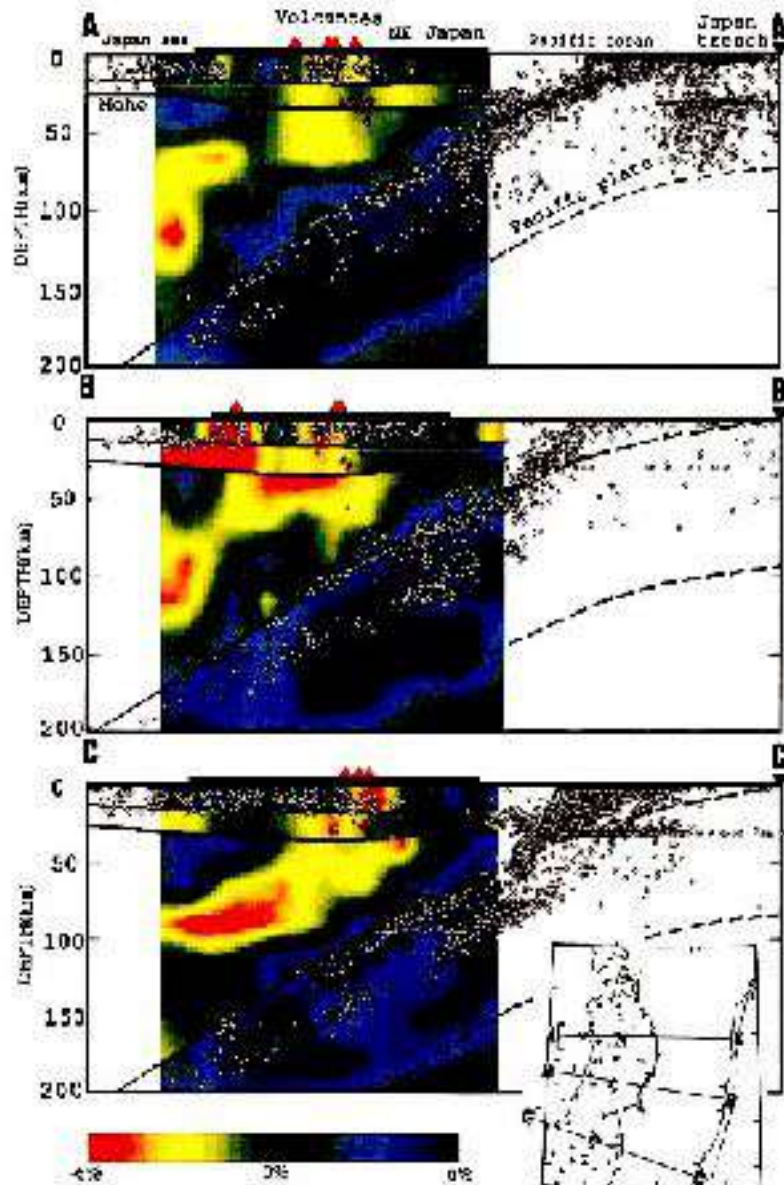
**Seismicity of Peru-Bolivia
Border Region: 1977 - 1997**



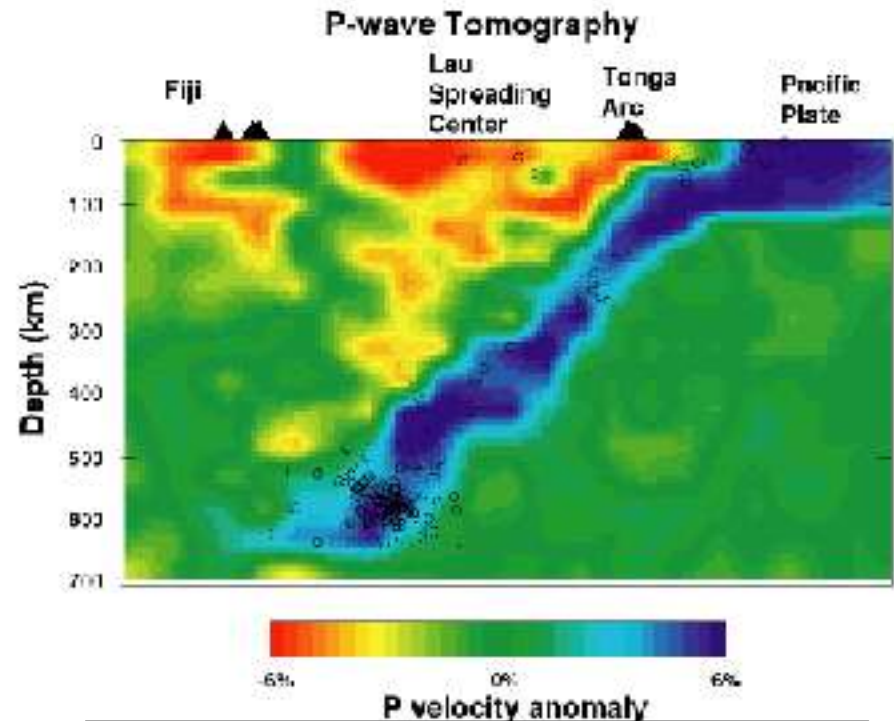
**Seismicity of Japan and Kuril Island:
1975 - 1995**

Earthquake and Plate Tectonics

Tomography (3D seismic)



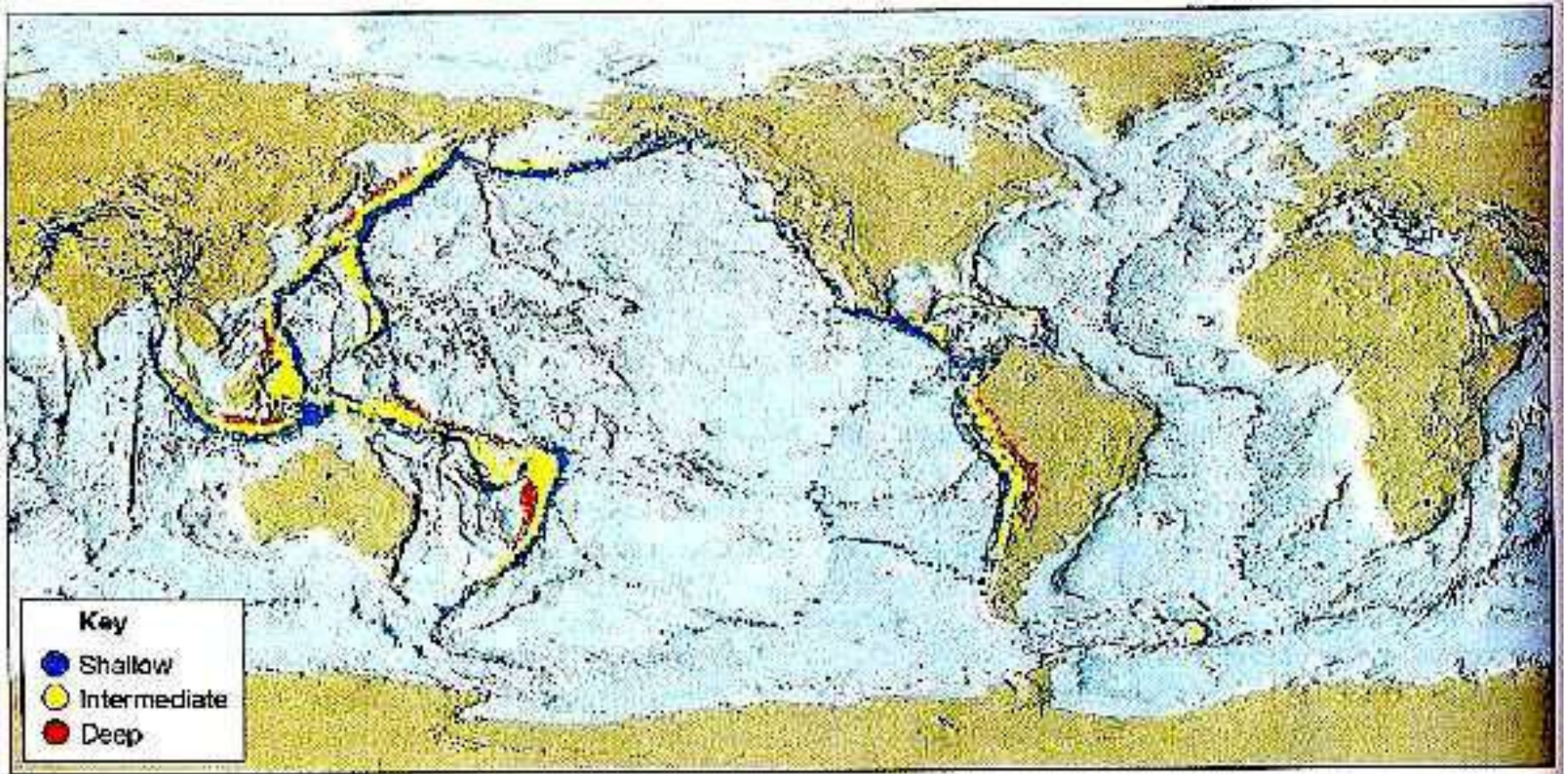
blue is fast...
interpreted as slab



note continuity of blue slab
to depths on order of 670 km

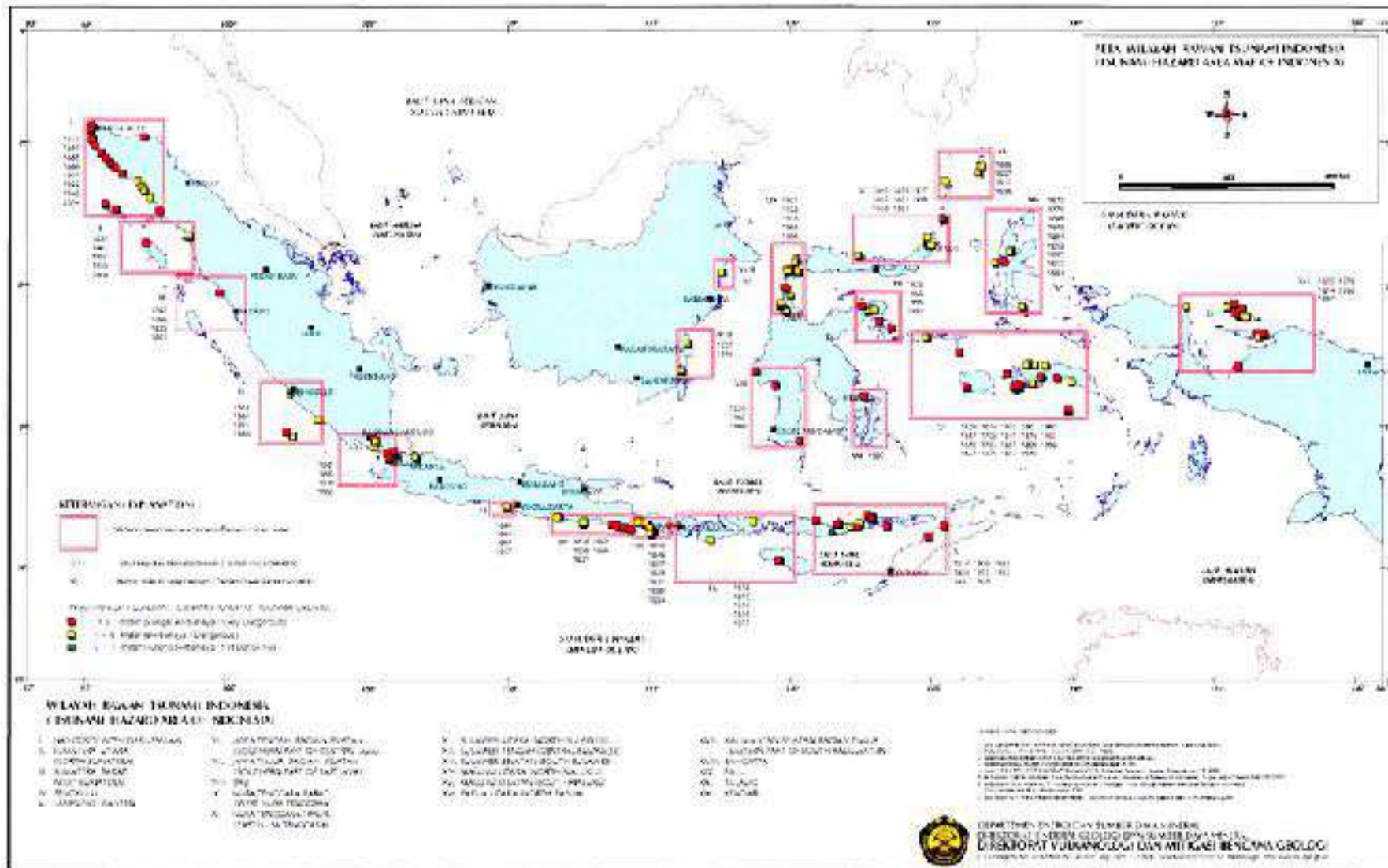
slab is cold and thus can have
earthquakes at greater depths

Earthquake and Plate Tectonics

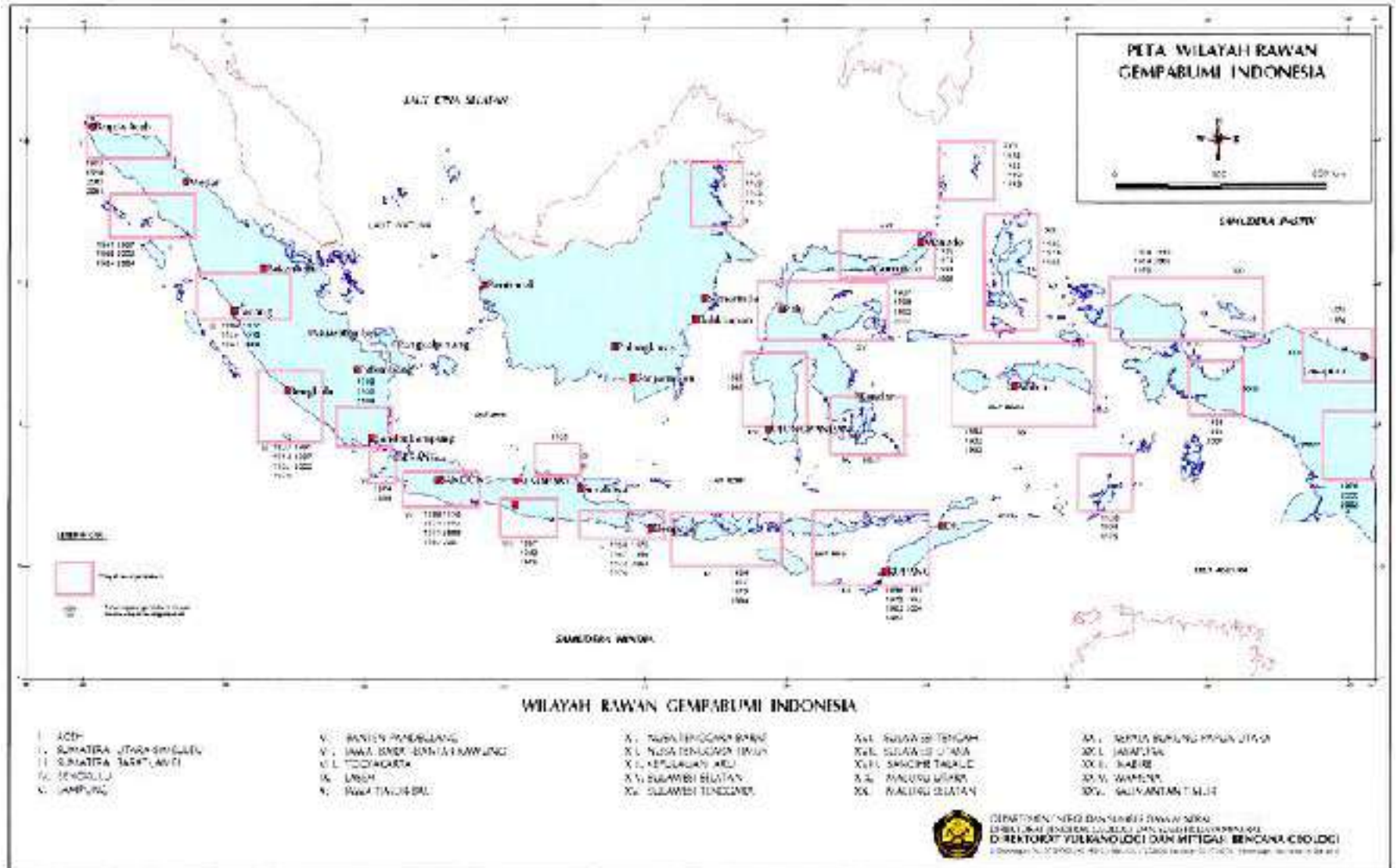


**Seismicity along subduction zones:
earthquakes are shallow, intermediate, and deep
but have systematic location related to subducting slab
shallow adjacent to trench and deep farthest away**

Tsunami Hazard Area Map of Indonesia



Earthquake Hazard Area Map of Indonesia



Map of Distribution of Active Faults and Destroying Earthquake Epicenters of Indonesia

